

Влияние изменения климата на вечную мерзлоту: наблюдения, моделирование, прогноз.

Impacts of changing climate on permaprost: observations and predictive modeling

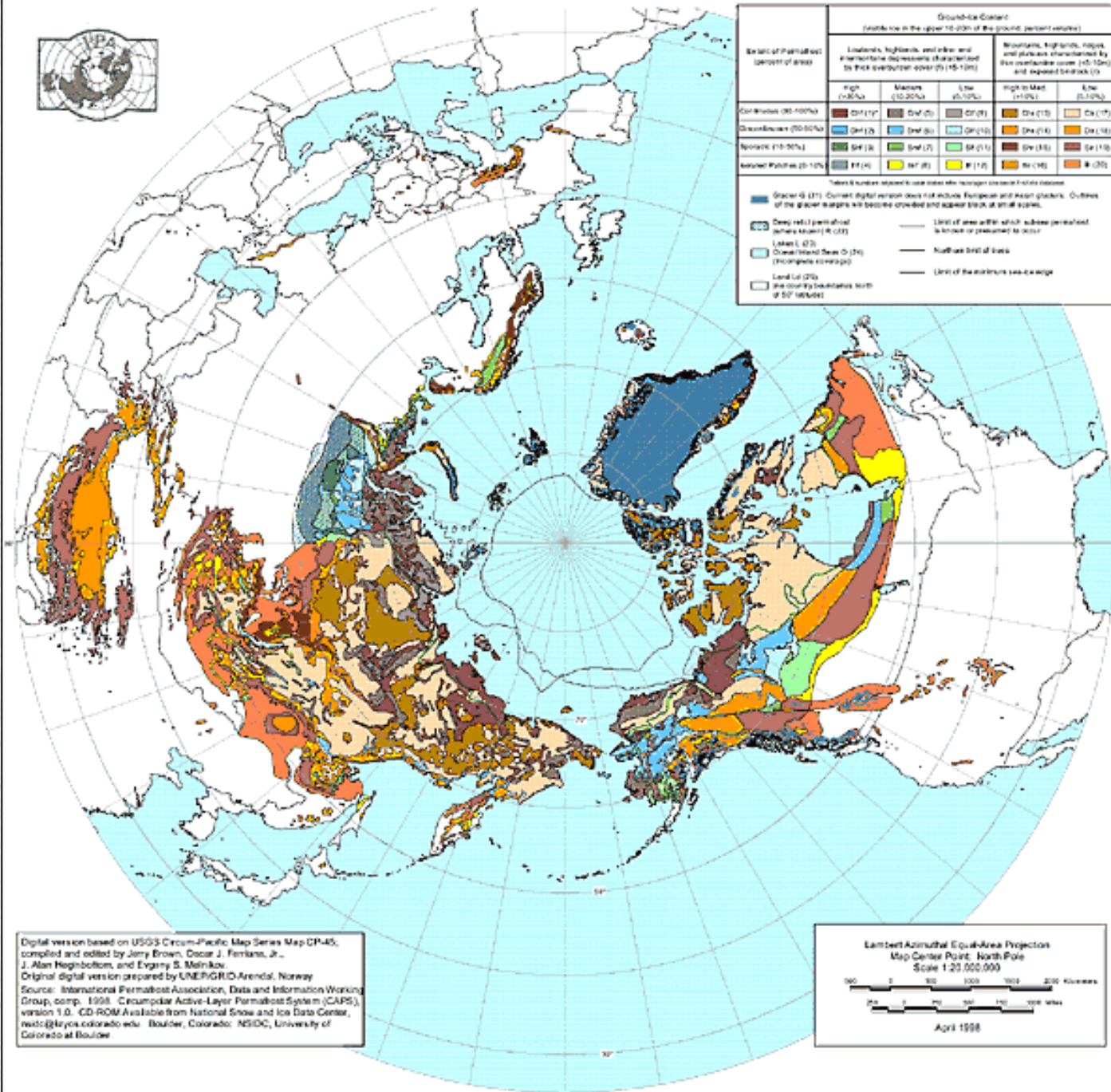
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CIRCUM-ARCTIC MAP OF PERMAFROST AND GROUND-ICE CONDITIONS



Extent of Permafrost (percent of area)	Ground-ice Content (weight ice in the upper 10-20% of the ground, percent water)				
	Lowlands, highlands, and other and flatter terrain depressions characterized by thick water-saturated cover (5-15%)			Mountains, highlands, ridges, and plateaus characterized by thin water-saturated cover (<5-10%) and exposed bedrock (3)	
	High (>30%)	Medium (10-30%)	Low (5-10%)	High in Mass (>10%)	Low (5-10%)
Continuous (80-100%)	CH (17)	CM (2)	CL (3)	CH (15)	CL (17)
Discontinuous (50-80%)	CH (2)	CM (3)	CL (14)	CH (14)	CL (14)
Sporadic (10-50%)	SH (3)	SM (2)	SL (1)	SH (14)	SL (15)
Seasonal Frozen (0-10%)	SH (4)	SM (5)	SL (12)	SH (14)	SL (20)

Notes: Numbers represent ice content in the average ice-rich terrain.

■ Glacier (G) (21) Current digital version does not include Fennoscandia and Asian glaciers. Outlines of the glacier margins will become crowded and appear thick at small scales.
■ Dry (rel.) permafrost (brown (brown) R) (22)
■ Layer 1 (23)
■ Down-land Sea 0 (24) (Incomplete coverage)
■ Land (L) (25) (No country boundaries north of 50° latitude)

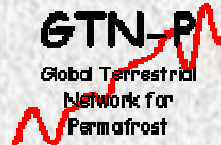
— Limit of area within which subsea permafrost is known or presumed to occur
— Northern limit of trees
— Limit of the maximum sea-ice coverage

Digital version based on USGS Circum-Pacific Map Series Map CP-45, compiled and edited by Jerry Brown, Oscar J. Ferrians, Jr., J. Alan Heginbottom, and Evgeniy S. Melnikov. Original digital version prepared by UNERGRID-Arendal, Norway. Source: International Permafrost Association, Data and Information Working Group, comp., 1998. Circumpolar Active-Layer Permafrost System (CAPS), version 1.0. CD-ROM Available from National Snow and Ice Data Center, nwtc@keyes.colorado.edu. Boulder, Colorado: NSIDC, University of Colorado at Boulder.

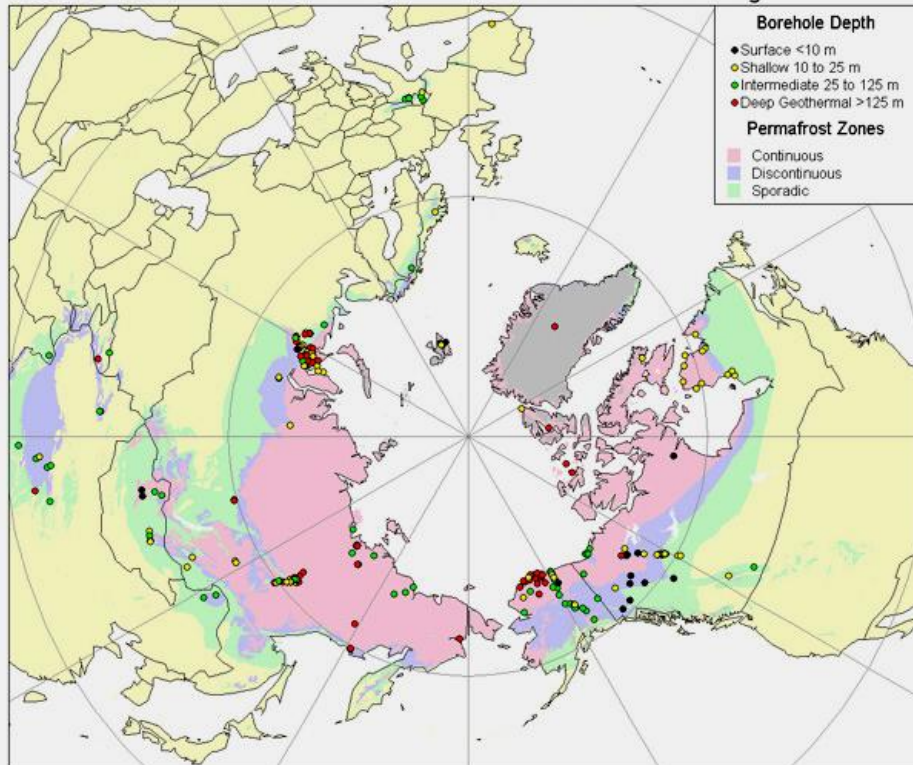
Lambert Azimuthal Equal-Area Projection
 Map Center Point: North Pole
 Scale 1:20,000,000

April 1998

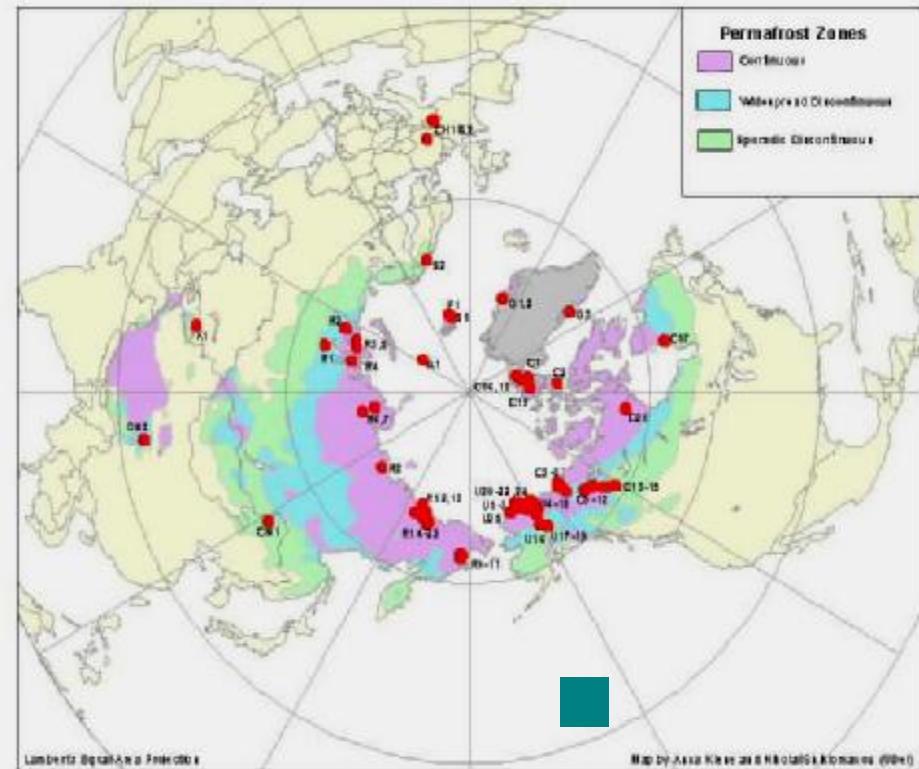
Permafrost observational programs



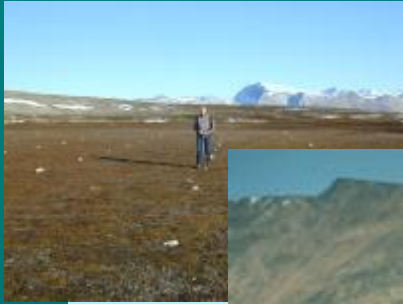
Candidate Boreholes for Permafrost Thermal Monitoring

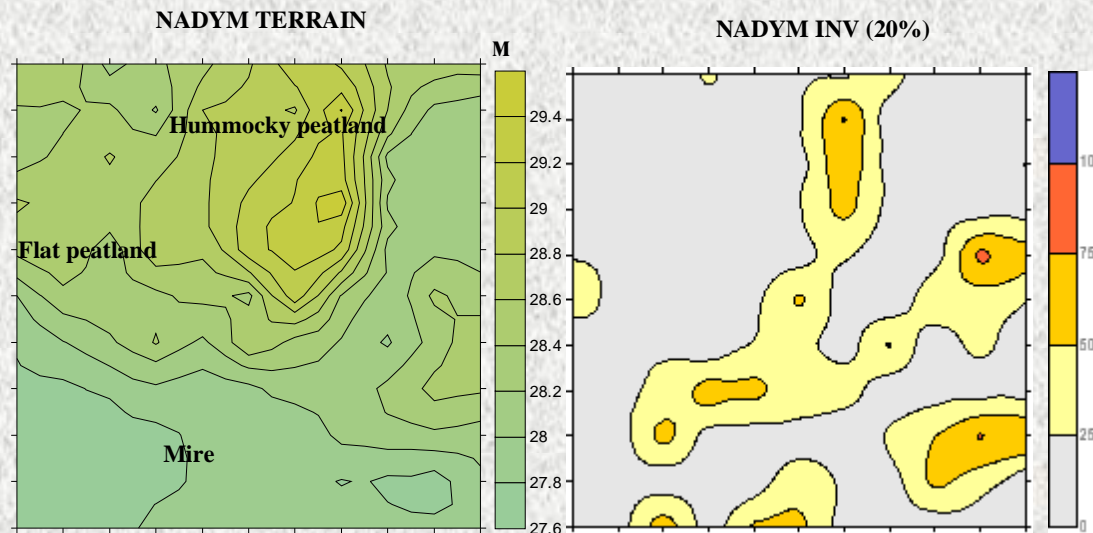
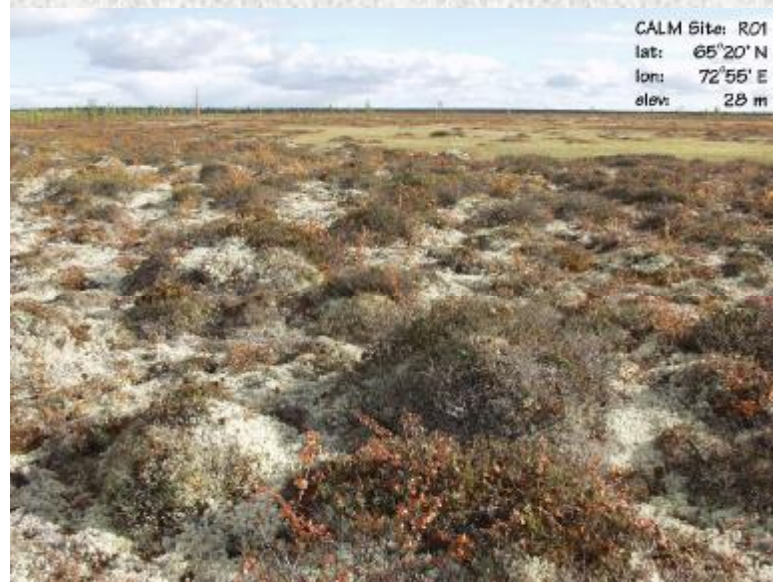
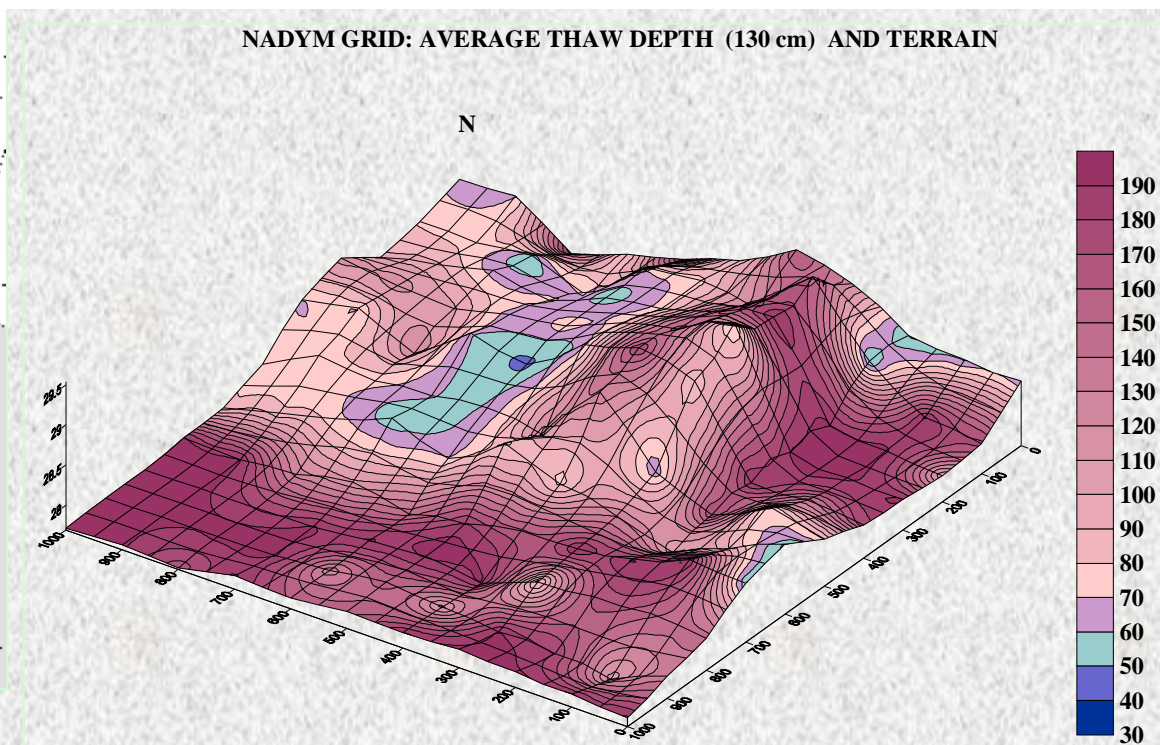


Circumpolar Active Layer Monitoring (CALM) Network



Permafrost measurements at CALM site on Svalbard

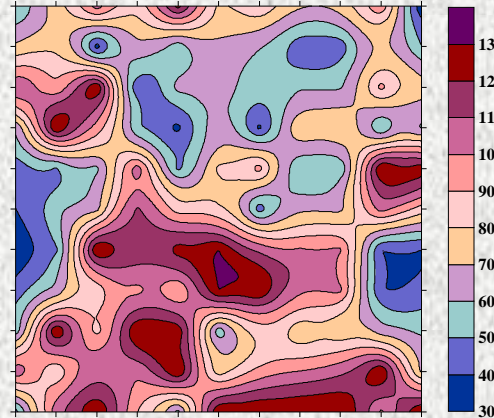




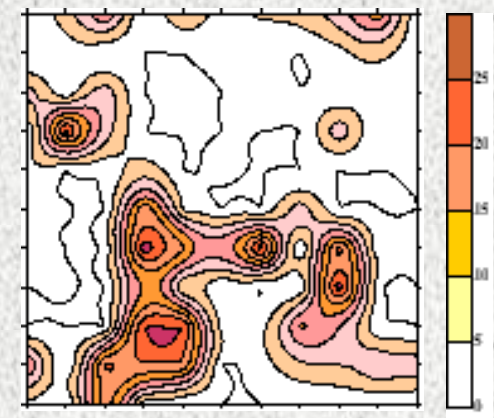
Vasiliev et al., 2005



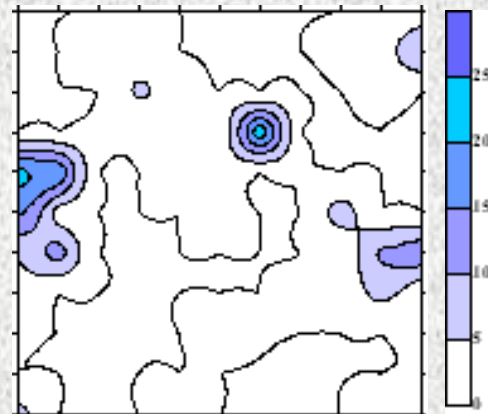
Active layer thickness, cm (1992)



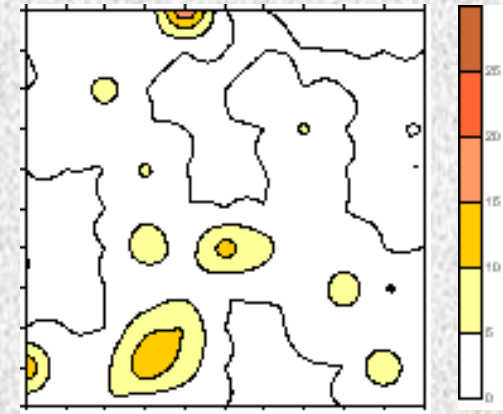
Alectoria (frequency, %)



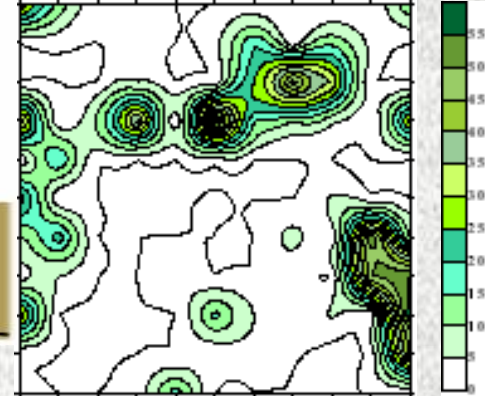
Rubus chamaemorus (frequency, %)



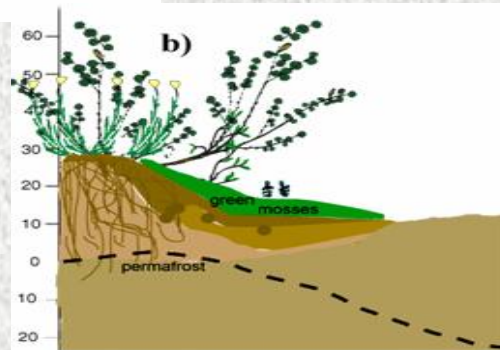
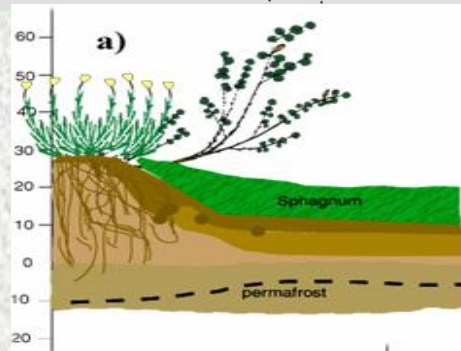
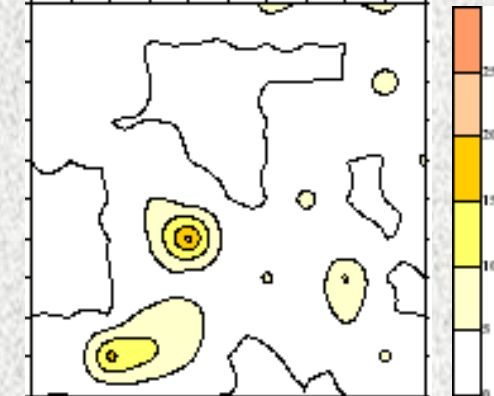
Sphaerophorus globosus (frequency, %)



Eriophorum vaginatum (frequency, %)



Cetraria nivalis (frequency, %)

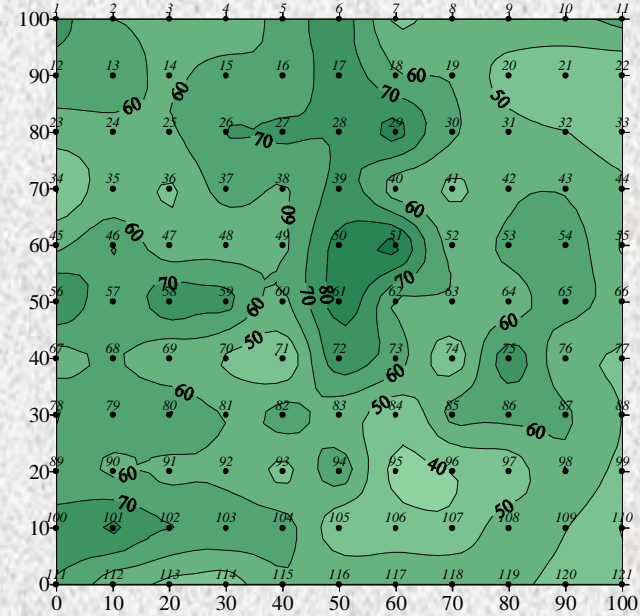


Malkova et al., 2005

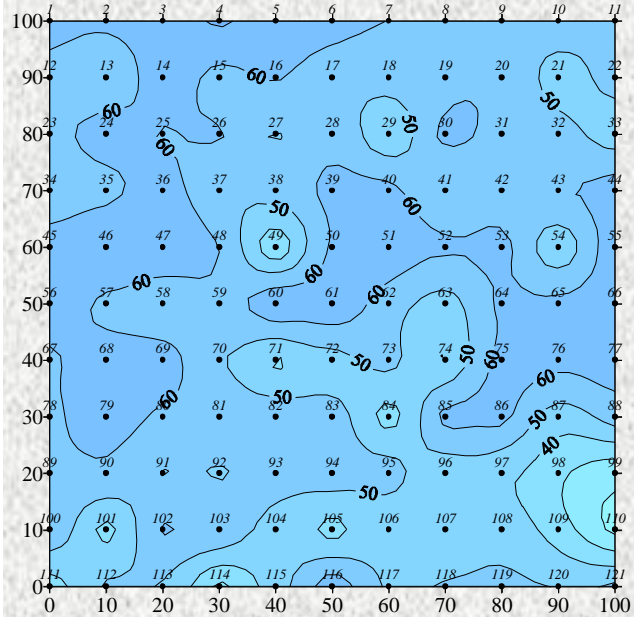


Cape Lavrentia

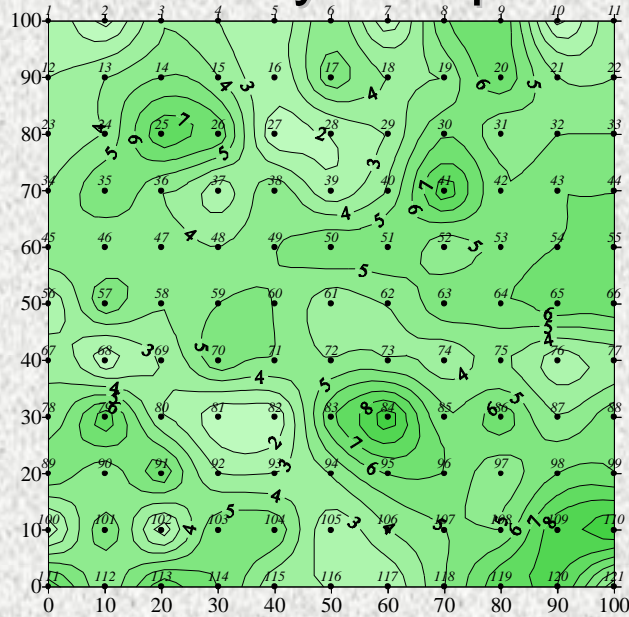
Thaw depth



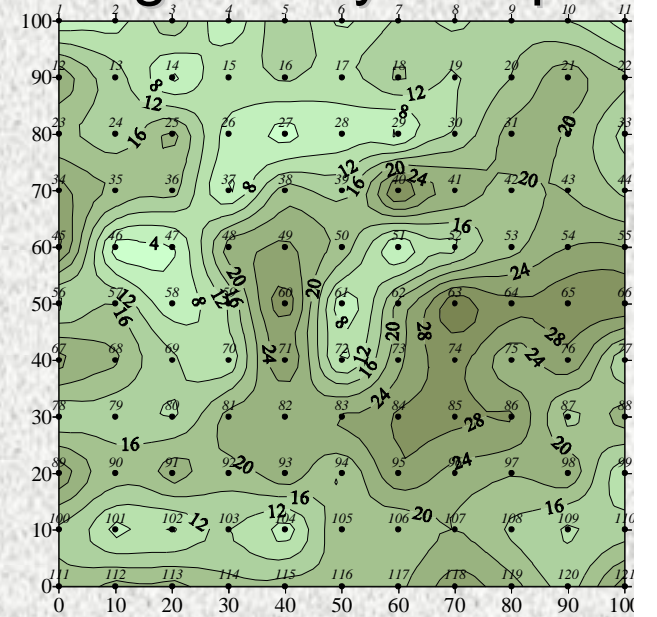
Soil moisture

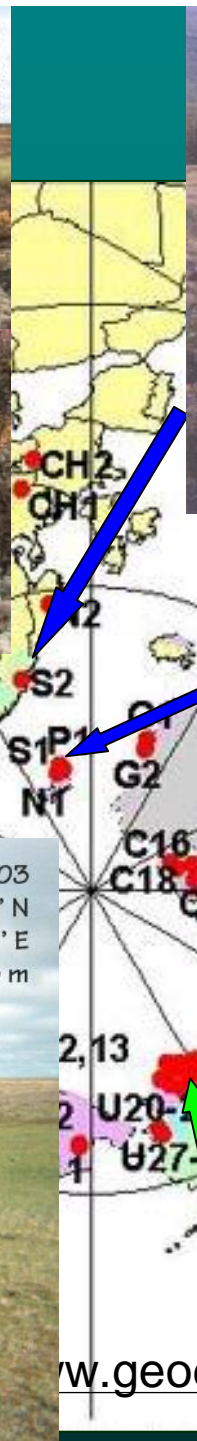
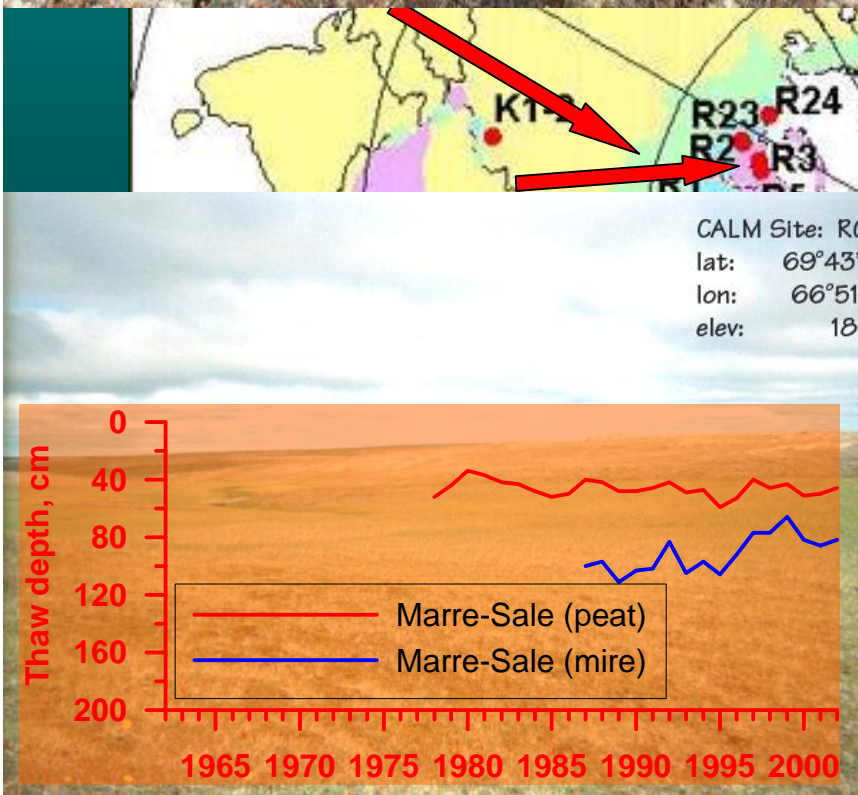
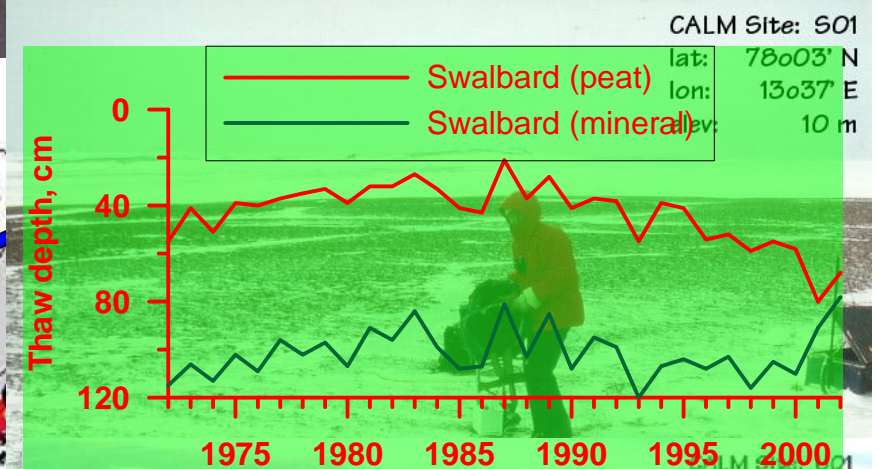
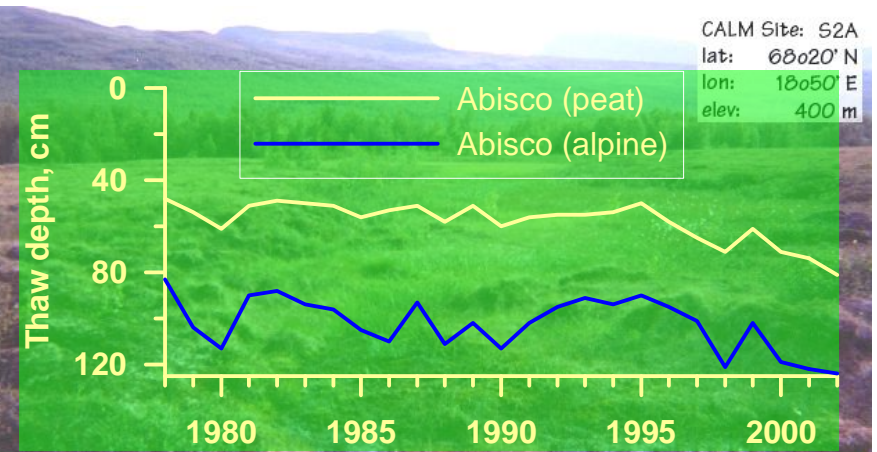
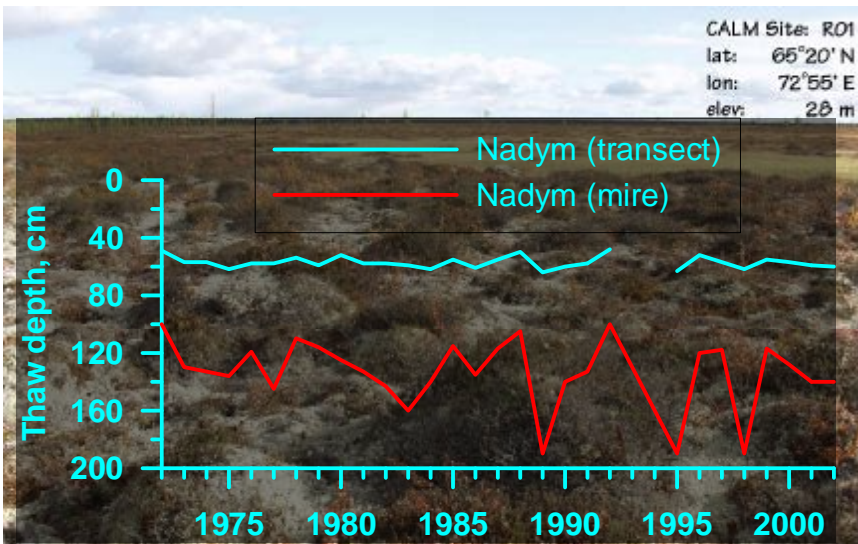


Moss layer depth

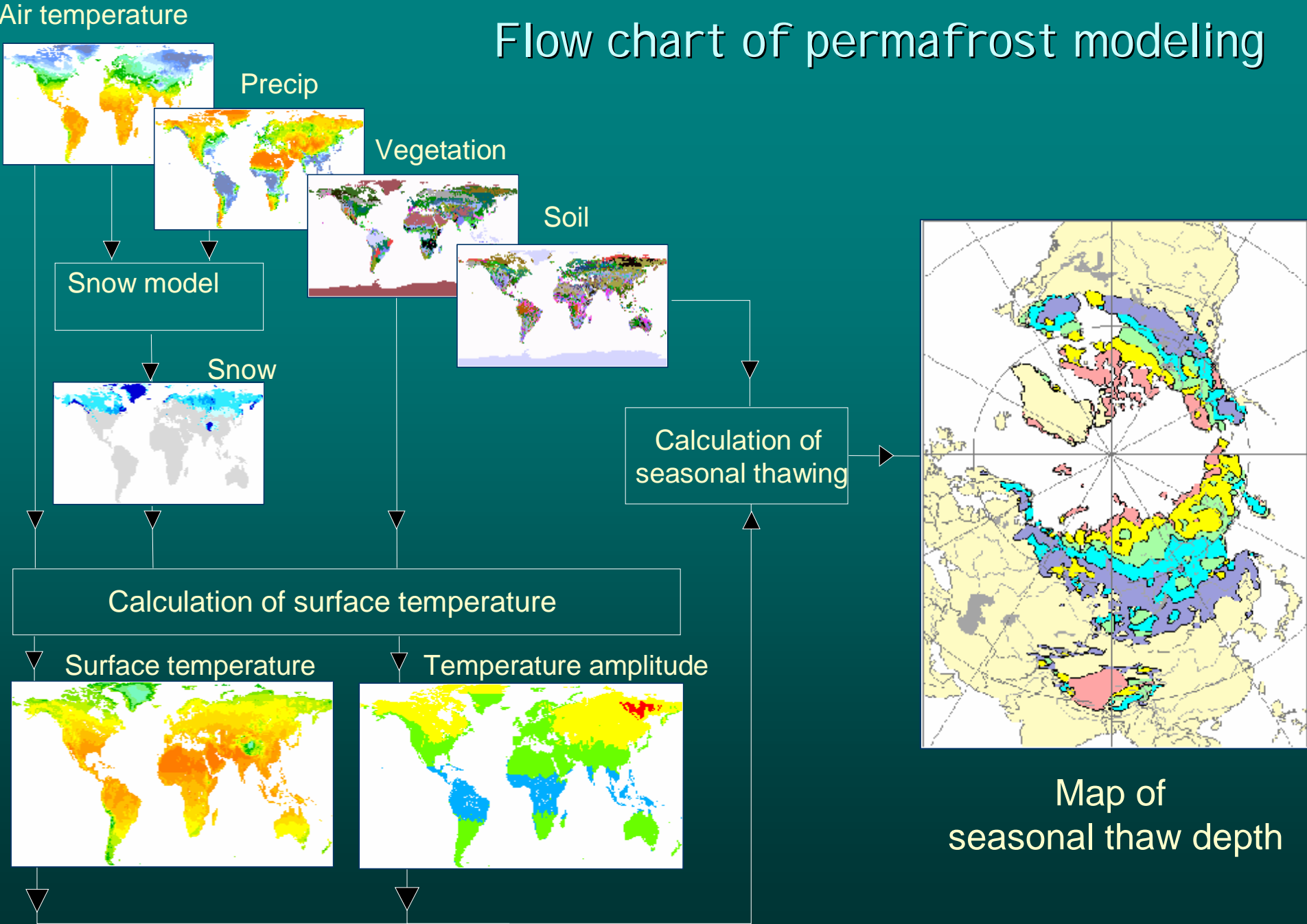


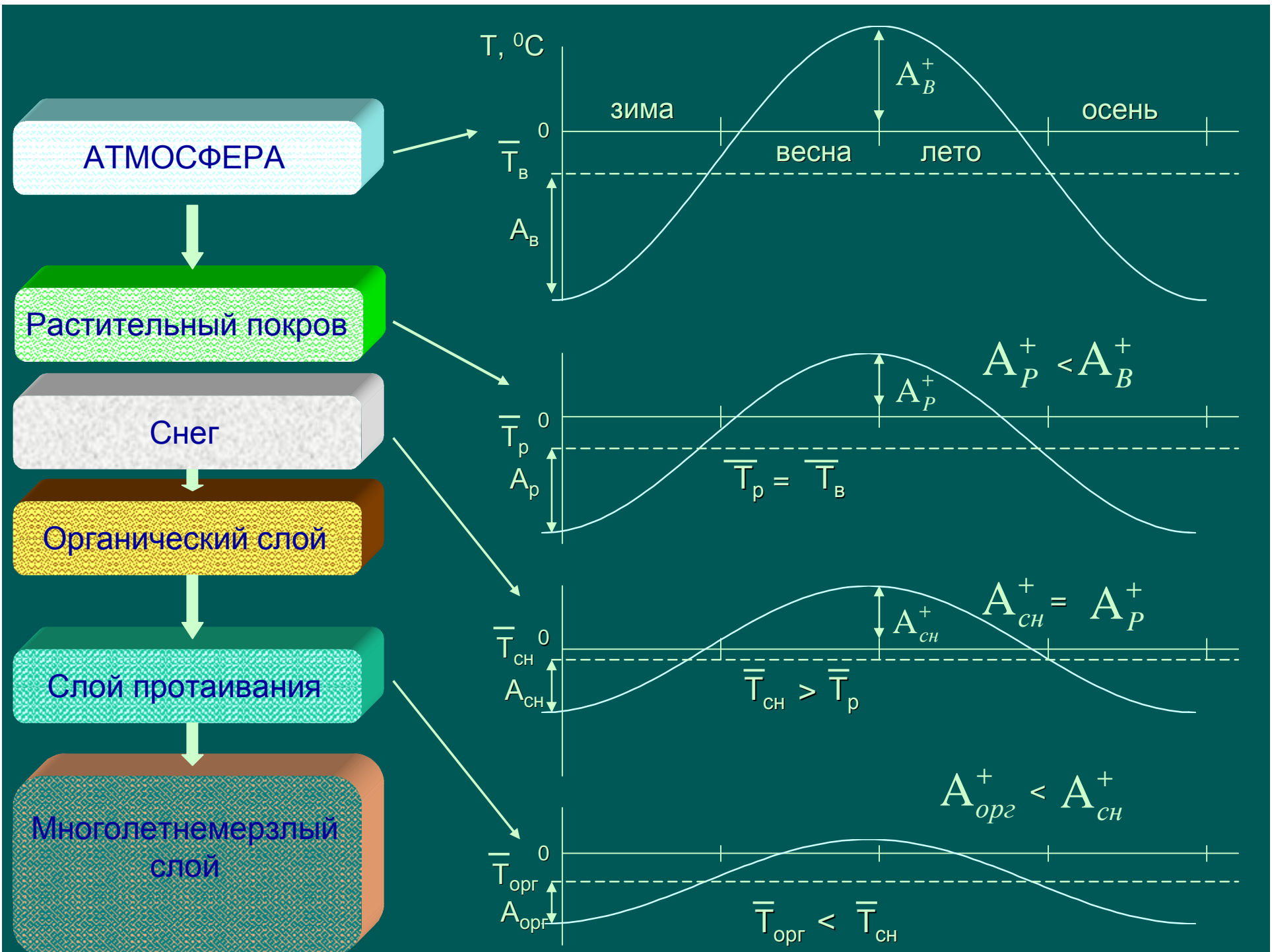
Organic layer depth





Flow chart of permafrost modeling





$$T_{\Pi} = T_B + \Delta T_{CH} + \Delta T_p, \quad A_{\Pi} = A_B - \Delta A_{CH} - \Delta A_p, \quad Z_{CH} = \sin^2 f \frac{1}{k} \left\{ \sum_{i=1}^k h_i (k - (i - 1)) \right\}$$

$$DT_{CH} = A_B \left(1 - \exp\left(-Z_{CH} \left(\frac{p \cdot c_{CH} \cdot r_{CH}}{P \cdot l_{CH}} \right)^{1/2} \right) \right) \quad DA_{CH} = A_B \left(1 - \exp\left(-Z_{CH} \left(\frac{p \cdot c_{CH} \cdot r_{CH}}{P \cdot l_{CH}} \right)^{1/2} \right) \right)$$

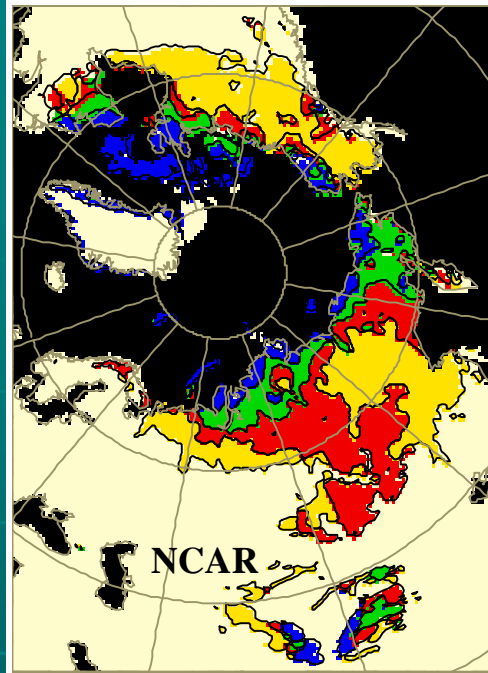
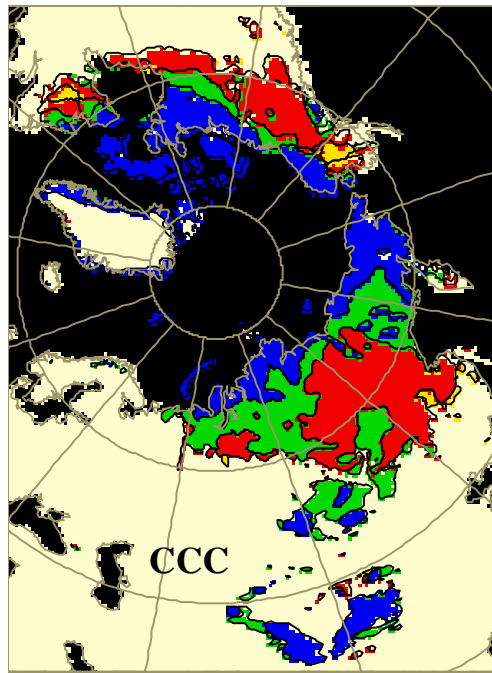
$$DT_p = \frac{DA_1 t_1 + DA_2 t_2}{P} \cdot \frac{2}{p} \quad DA_p = \frac{DA_1 t_1 + DA_2 t_2}{P}$$

$$DA_1 = (A_p - T_p) \left(1 - \exp\left(-Z_p \left(\frac{p}{K_p^- \cdot 2t_1} \right)^{1/2} \right) \right) \quad DA_2 = (A_p + T_p) \left(1 - \exp\left(-Z_p \left(\frac{p}{K_p^+ \cdot 2t_2} \right)^{1/2} \right) \right)$$

$$T_p = T_B + DT_{CH},$$

$$A_p = A_B - \Delta A_{CH}, \quad T_M = \frac{0.5 \cdot T_{\Pi} \cdot (l_M + l_T) + A_{\Pi} \frac{l_T - l_M}{p} \left(\frac{T_{\Pi}}{A_{\Pi}} \arcsin \frac{T_{\Pi}}{A_{\Pi}} + \left(1 - \frac{T_{\Pi}^2}{A_{\Pi}^2} \right)^{1/2} \right)}{l^*}$$

$$Z = \frac{2(A_{\Pi} - T_M) \cdot \left(\frac{l \cdot P \cdot C}{p} \right)^{1/2} + \frac{(2A_M \cdot C \cdot Z_M + Q_{\phi} \cdot Z) \cdot Q_{\phi} \cdot \left(\frac{l \cdot P \cdot C}{p \cdot C} \right)^{1/2}}{2A_M \cdot C \cdot Z_M + Q_{\phi} \cdot Z + (2A_M \cdot C + Q_{\phi}) \cdot \left(\frac{l \cdot P \cdot C}{p \cdot C} \right)^{1/2}}{2A_M \cdot C + Q_{\phi}}$$



Projected for 2050 changes of seasonal thaw depth (relative to 2000)

0 – ocean

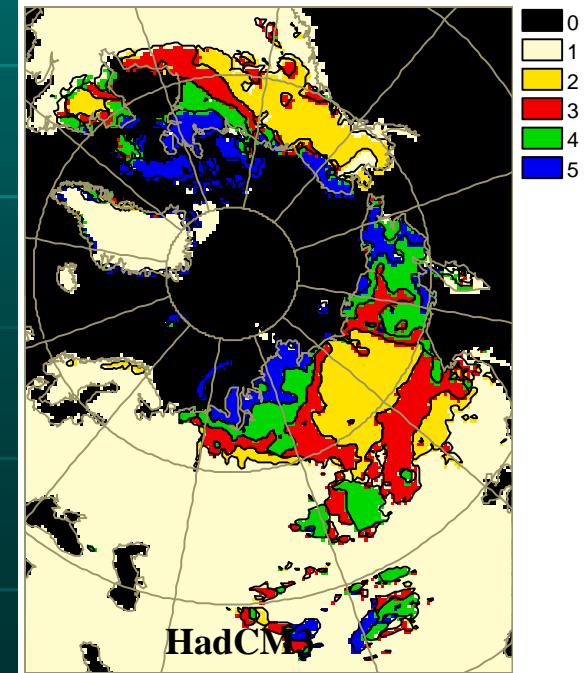
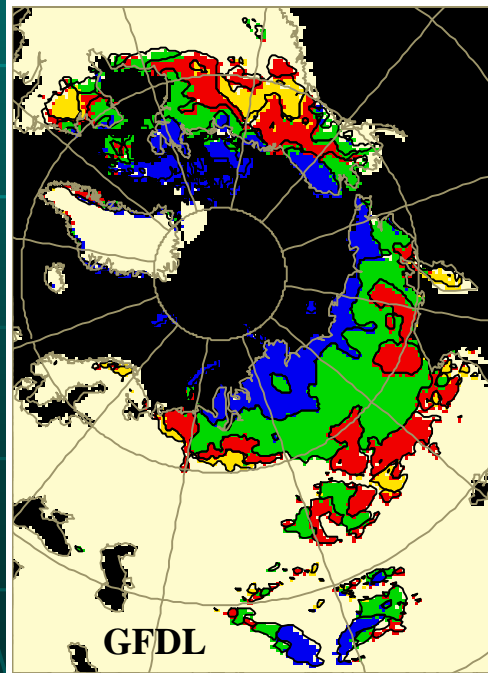
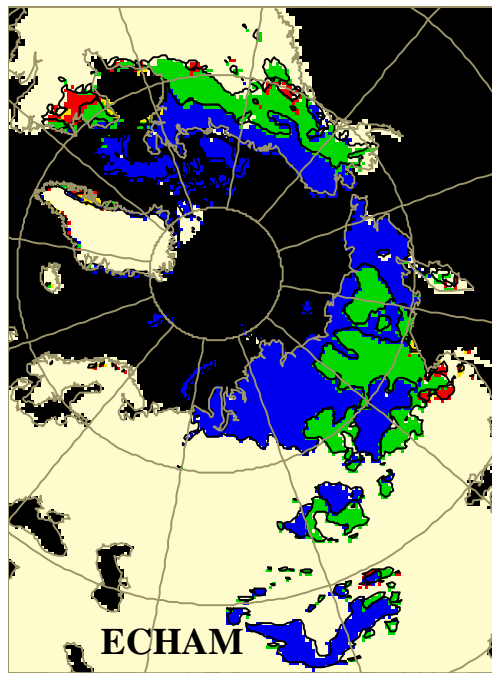
1 – permafrost-free land

2 – 0% - 20% increase

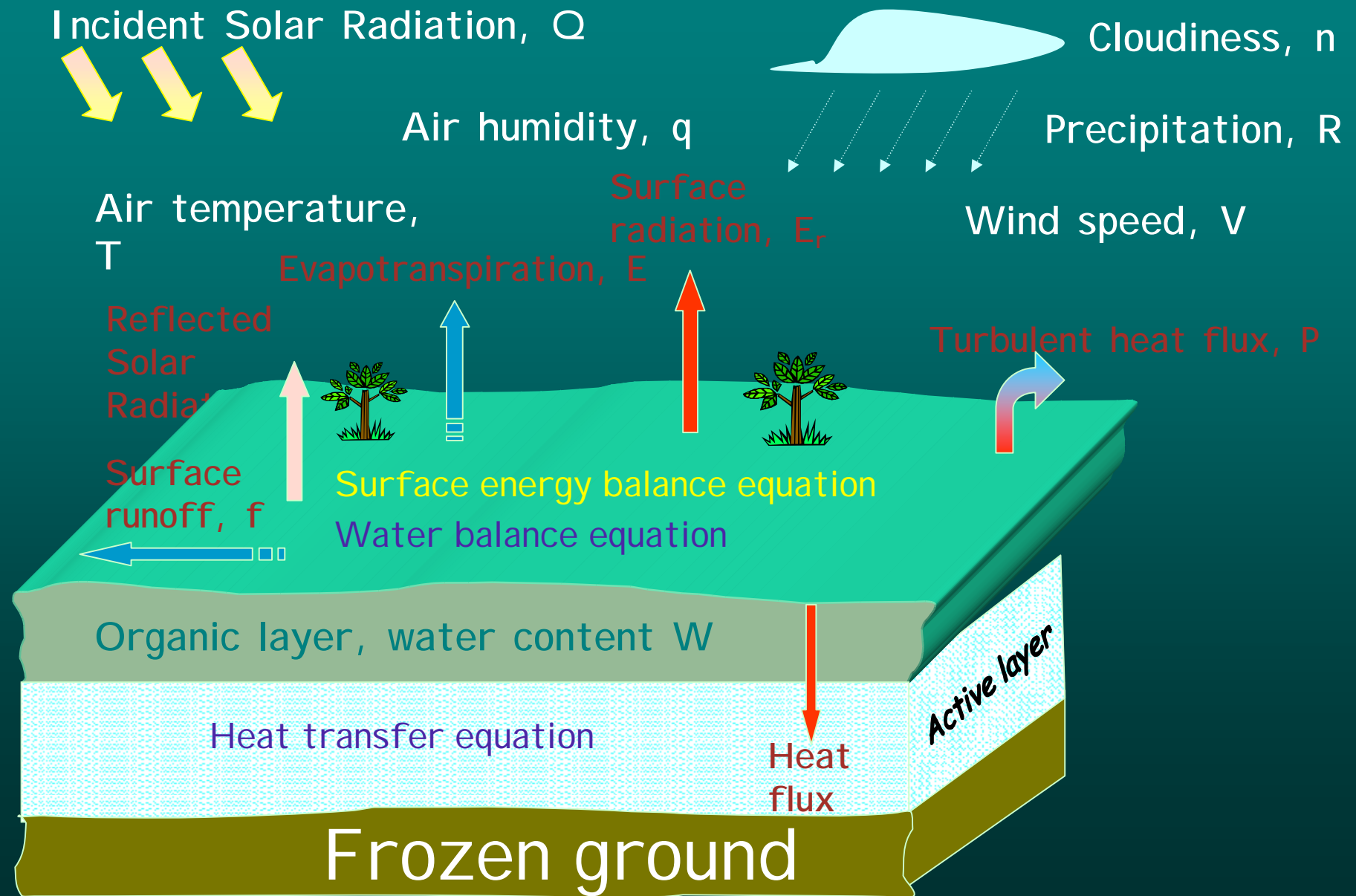
3 – 20% - 30% increase

4 – 30%-50% increase

5 – >50% increase



Dynamical permafrost modeling



Surface energy balance equation

$$Q(1-\alpha) - E_r(T_s) - P(T_s) - B(T_s) - LE(T_s) = 0$$

$$E_r(T_s) = E_r^o(T_s) (1 - 0.79 n) + 4 d_s T^3(T_s - T); \quad E_r^o(T_s) = d_s T^4(0.39 - 0.058 e^{1/2});$$

$$P(T_s) = r_a c_a v D_t (T_s - T); \quad B = I \frac{dT}{dz} \quad \begin{array}{l} E = E_o, \quad \text{if } w > w_k \\ E = E_o w/w_k, \quad \text{if } w \leq w_k \end{array} \quad E_o = r_a v D_t (e_u - e)$$

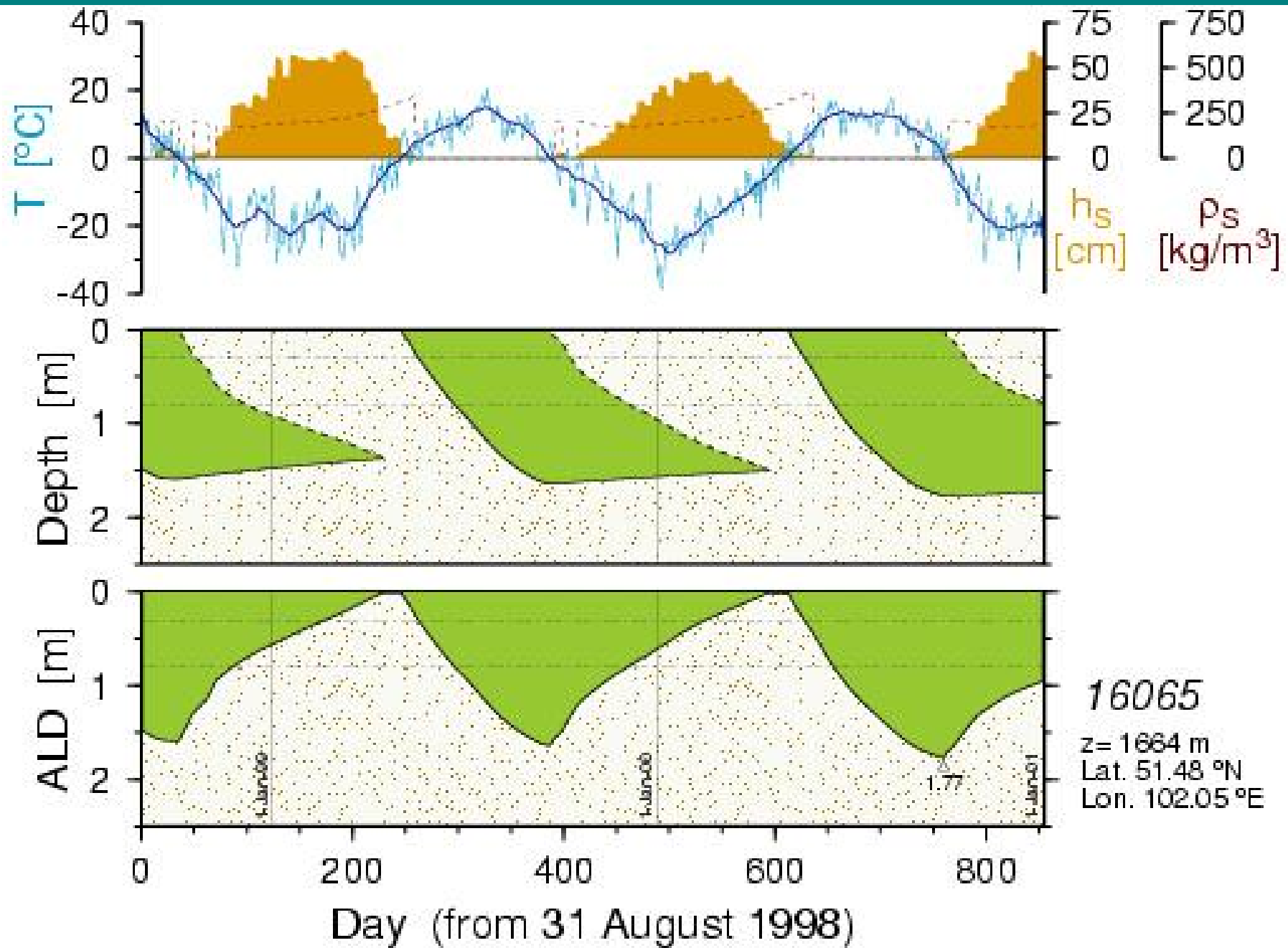
Heat transfer equation

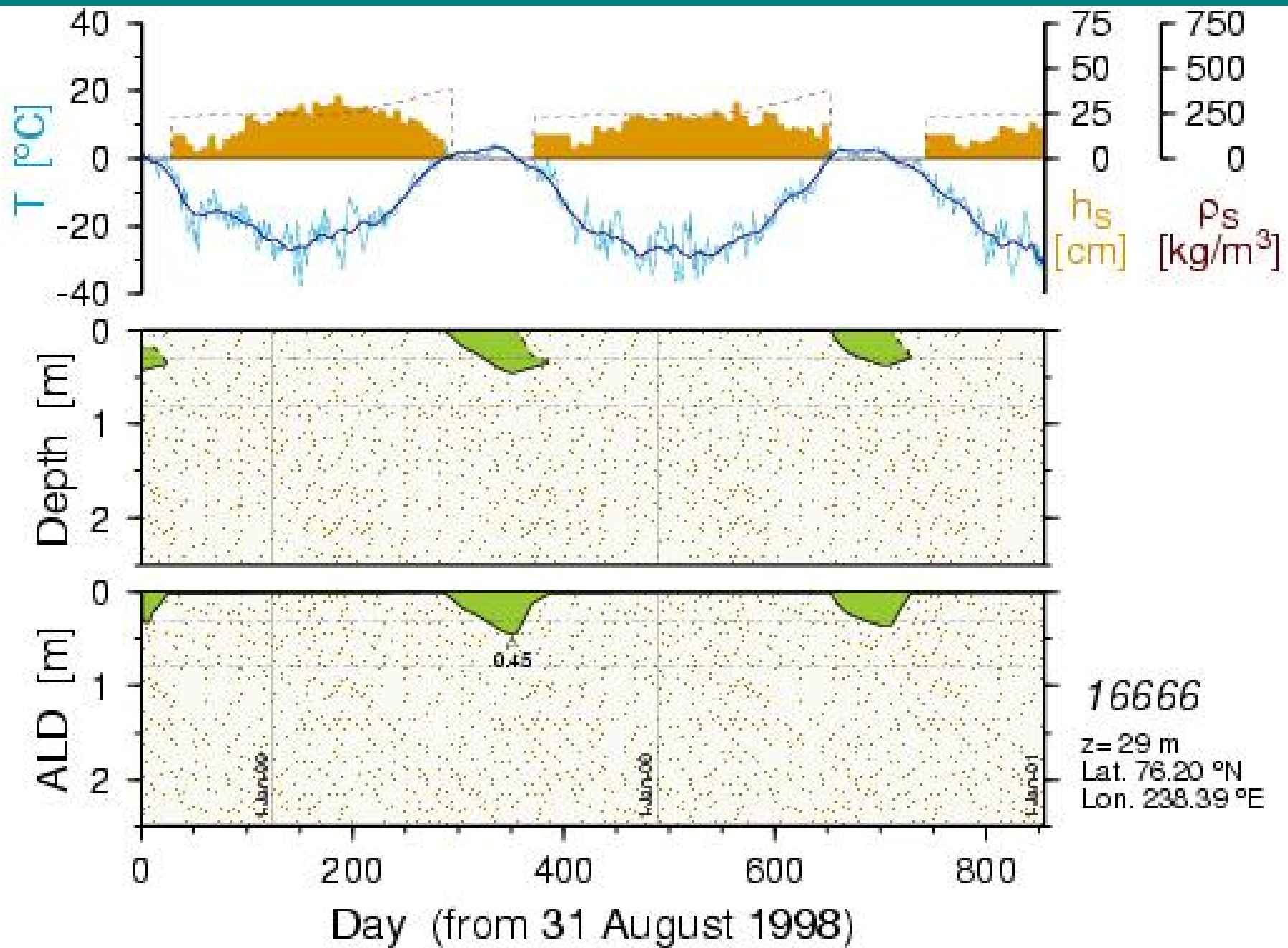
$$rc \frac{dT}{dt} = \frac{d}{dz} \left(I \frac{dT}{dz} \right) \quad \frac{dz_j}{dt} = (-1)^{j+1} \frac{1}{wL} \left(I \frac{dT}{dz} \Big|_{z=z_{j+0}} - I \frac{dT}{dz} \Big|_{z=z_{j-0}} \right)$$

Water balance equation

$$z_t \frac{dw}{dt} = R - E - f \quad f = R \frac{w}{w_k} \sqrt{m^2 \left[1 - \left(1 - \frac{E_o}{R} \right)^2 \right] + \left(1 - \frac{E_o}{R} \right)^2}, \quad \text{if } R > E_o$$

$$f = m R w/w, \quad \text{if } R \leq E_o$$





Stochastic permafrost modeling

Stefan Solution for Depth of Thaw:

$$z = \sqrt{\frac{2ks(nDDT)}{r\rho L}}$$

z = active-layer thickness

k = thermal conductivity of soil

s = unit conversion factor

n = n-factor

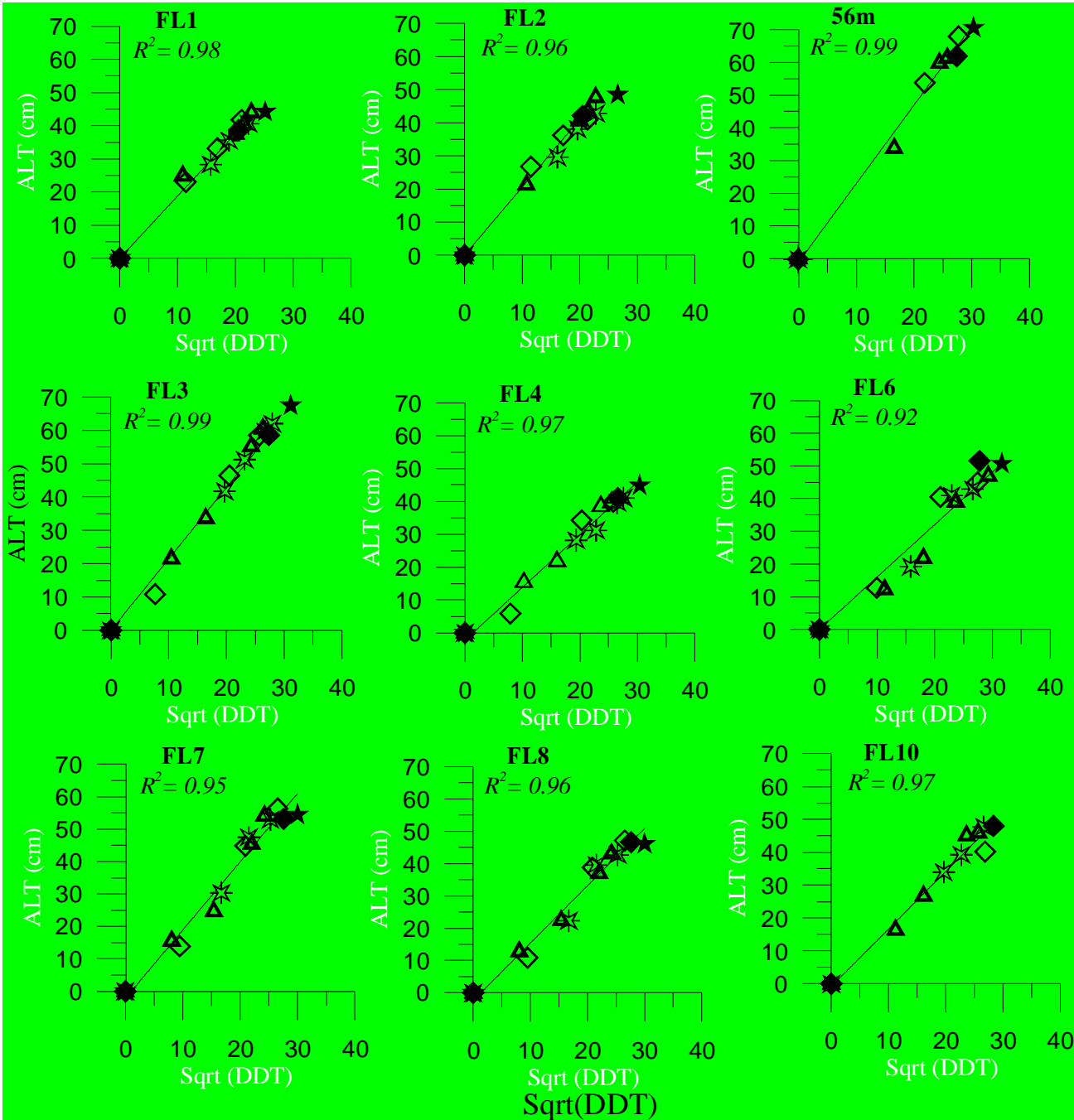
DDT = Degree Days of Thawing

ρ = soil density

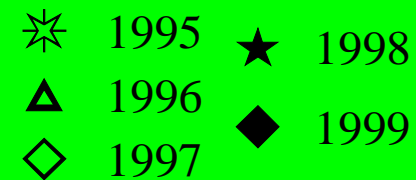
w = soil moisture

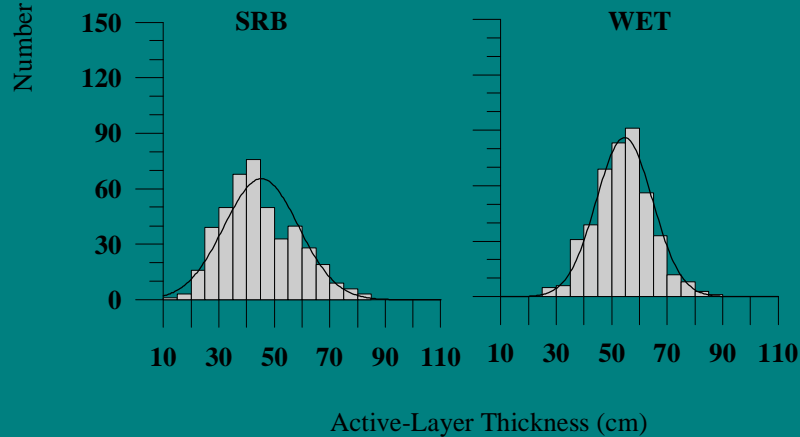
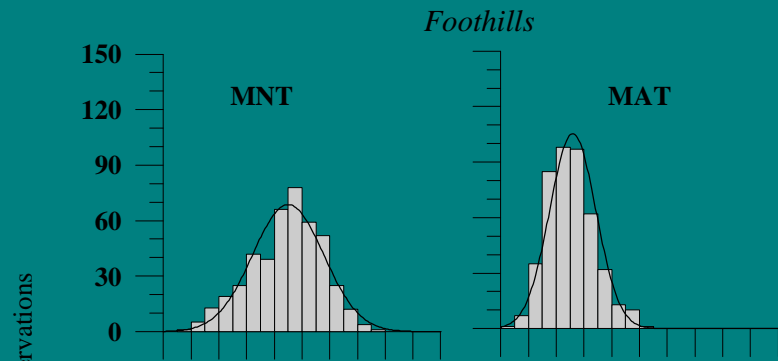
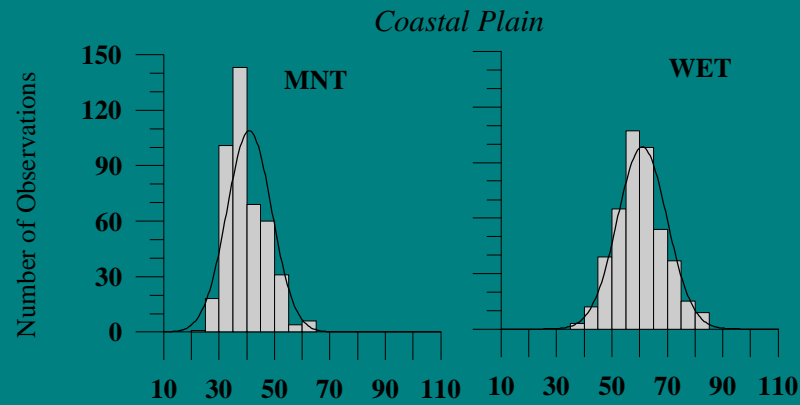
L = latent heat of fusion

$$z = E \cdot \sqrt{DDT}$$



Site	Land-Cover Class	<i>E</i>
<i>Coastal Plain</i>		
FL-1	Moist Nonacidic	1.87
FL-2	Wet Nonacidic	1.97
56m	Wet Tundra	2.42
<i>Foothills</i>		
FL-3	Moist Nonacidic	2.29
FL-4	Moist Acidic	1.62
FL-6	Moist Acidic / Water Track	1.71
FL-7	Wet Acidic	2.08
FL-8	Moist Acidic	1.65
FL-10	Shrubs / Water Track complex	1.73





Conventional formalism:

$$z = E C, \quad (1)$$

$$\text{where } E = [(2nl)/(rwL)]^{1/2}, \quad C = DDT^{1/2}$$

Stochastic formalism:

Variables are represented as mean value \bar{f} and fluctuation $f\hat{c}$:

$$f = \bar{f} + f\hat{c} \quad (2)$$

Stefan solution becomes

$$\bar{z} + z\hat{c} = \bar{E} \bar{C} + E\hat{c} \bar{C} + \bar{E} C\hat{c} + E\hat{c} C\hat{c} \quad (3)$$

Averaging of (3) yields equation for the mean ALT:

$$\bar{z} = \bar{E} \bar{C} \quad (4)$$

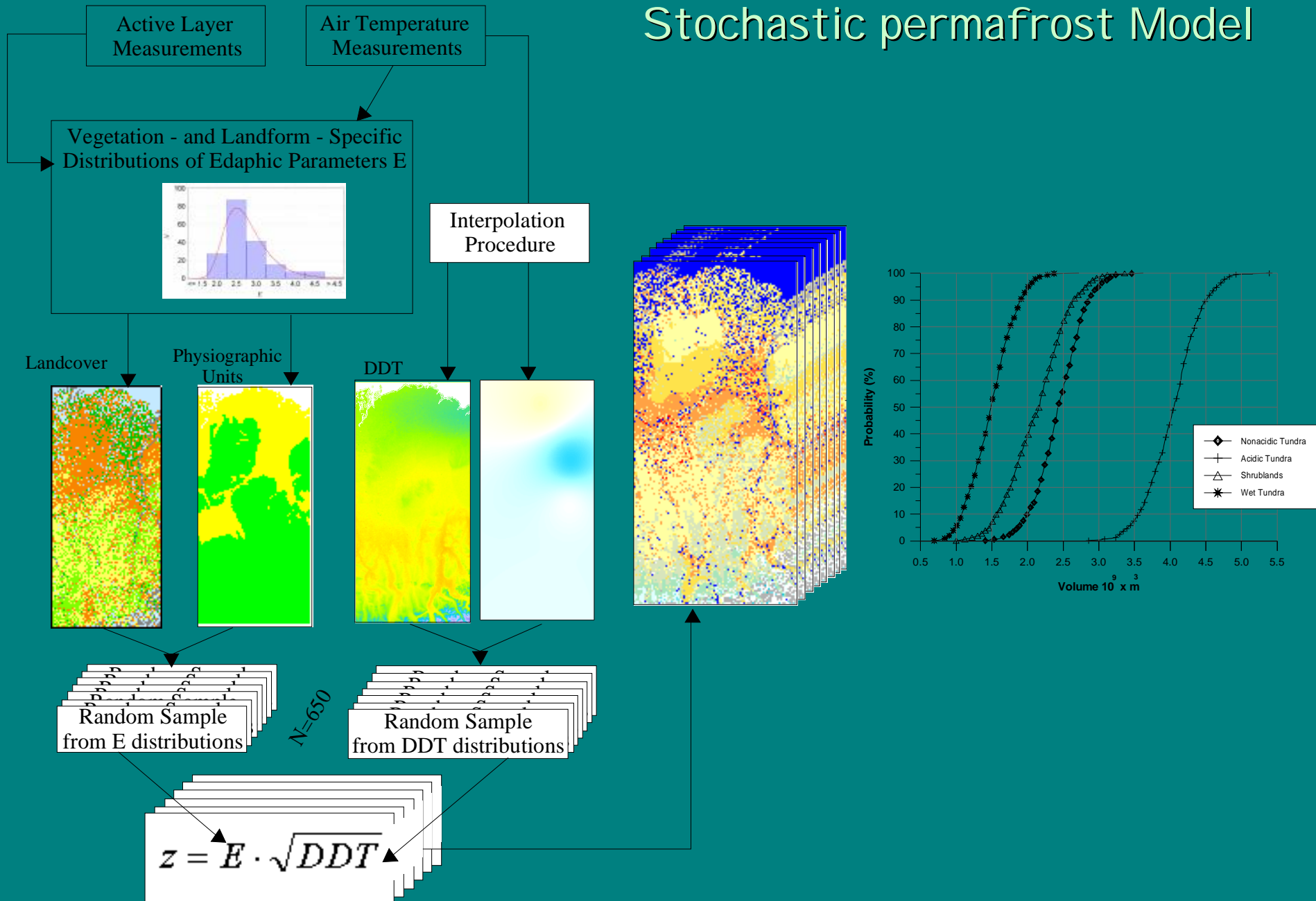
(4) - (3) yields equation for fluctuation:

$$z\hat{c} = E\hat{c} \bar{C} + \bar{E} C\hat{c} + E\hat{c} C\hat{c} \quad (5)$$

Square of (6) yields equation for variance:

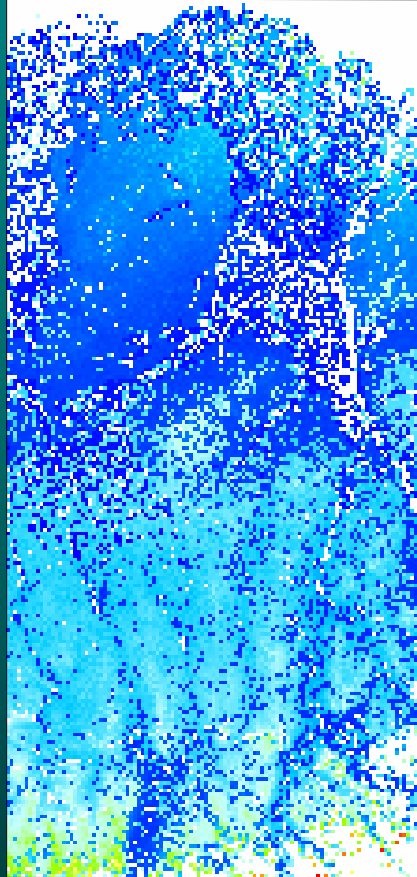
$$\overline{z^2} = \overline{E^2 C^2} + 2\overline{E\hat{c} E C} + 2\overline{E^2 C\hat{c}} + \overline{E^2 C^2} + 2\overline{E\hat{c} E C^2} + \overline{E^2 C^2}$$

Stochastic permafrost Model

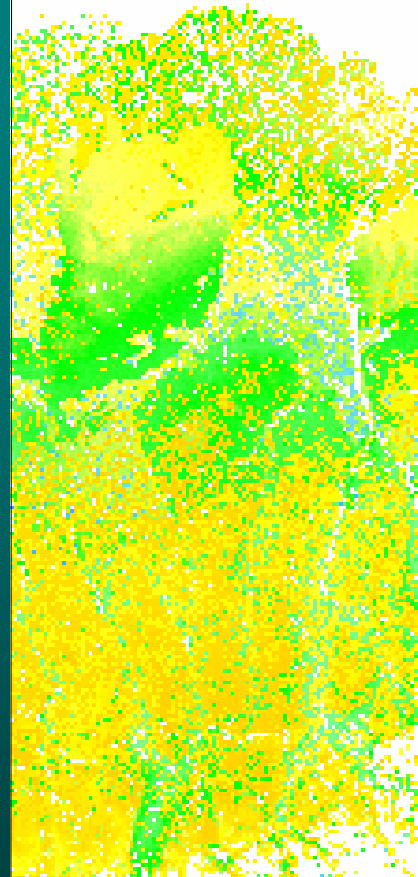


Probabilities of the active-layer thickness

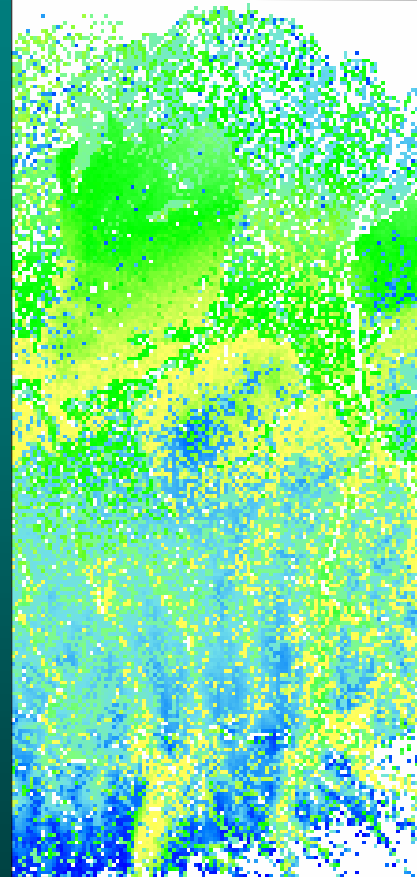
20-40 cm



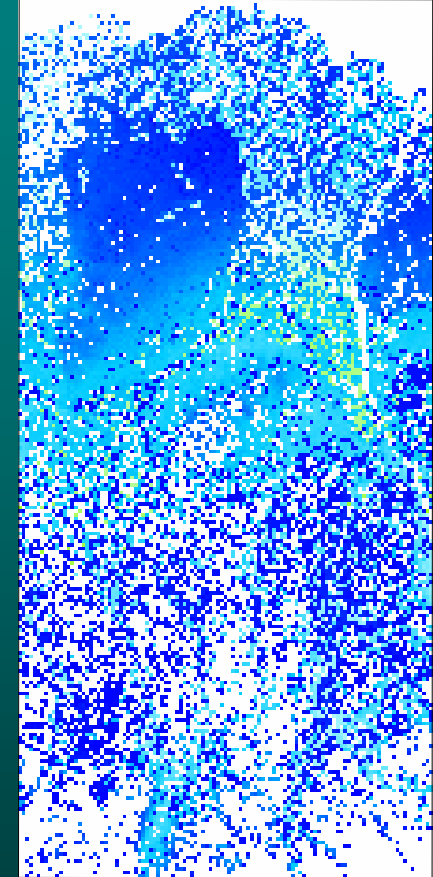
40-60 cm



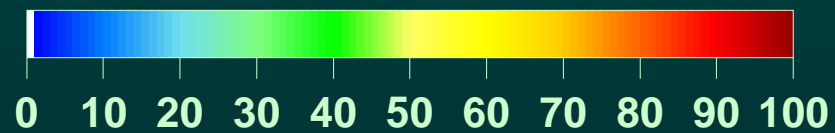
60-80 cm



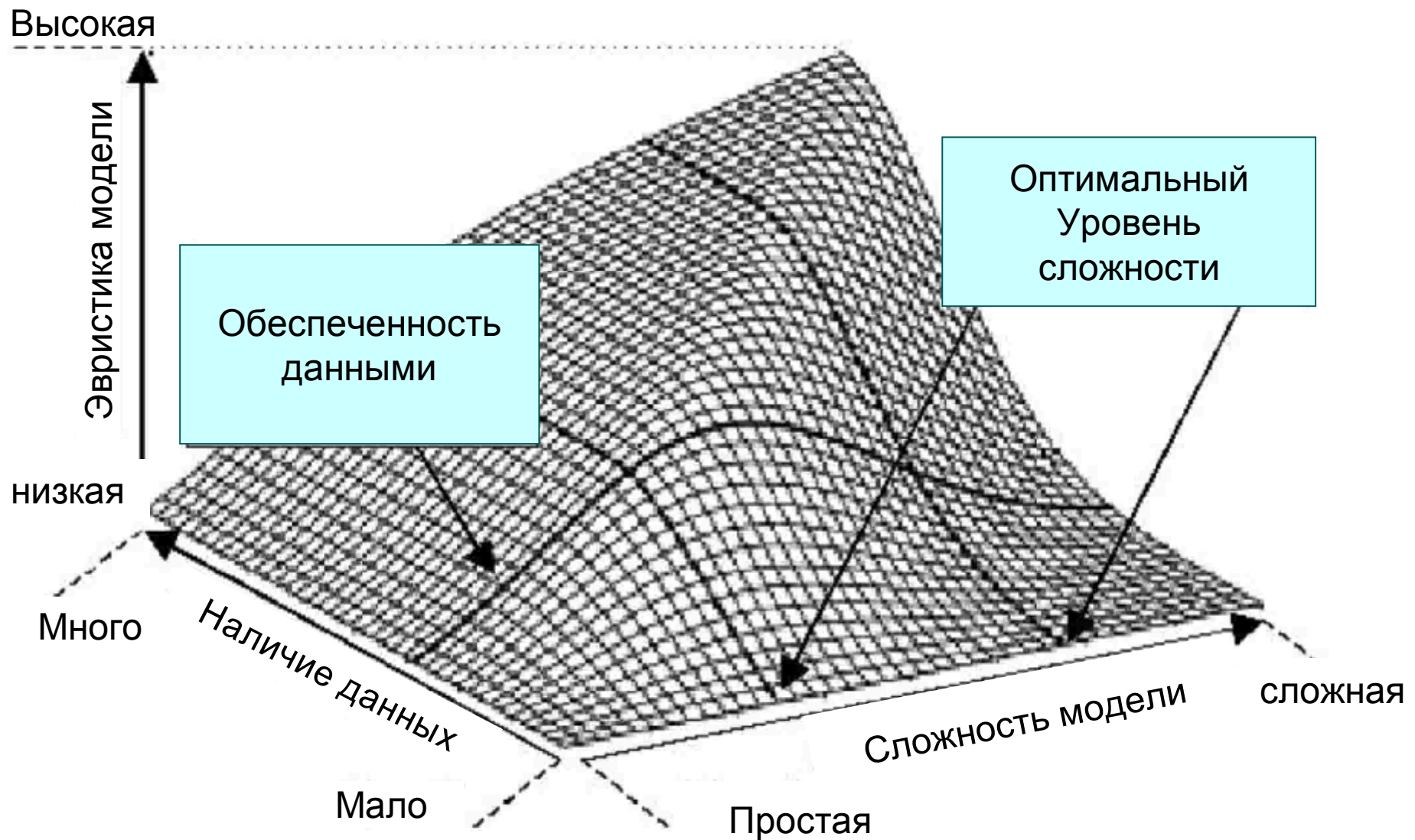
>80 cm



Probabilities (%)



Выбор оптимального уровня сложности модели в зависимости от имеющихся данных



Typical permafrost terrain



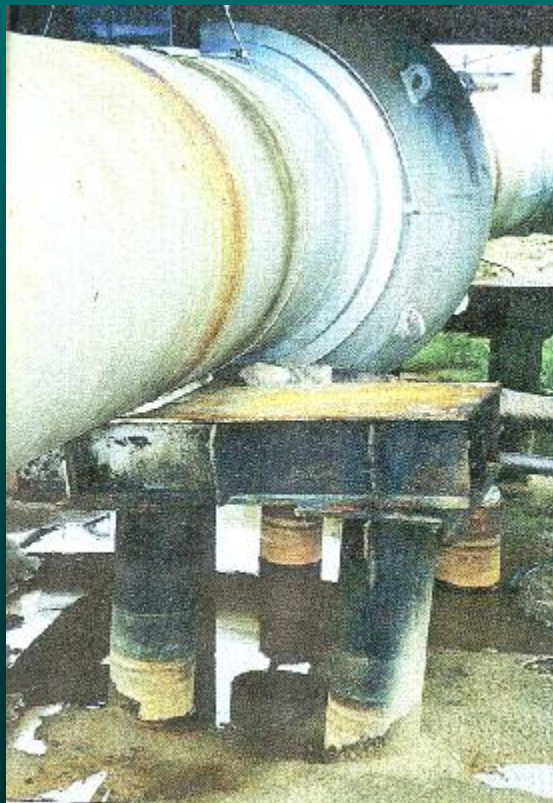
Photo by N.Shiklomanov

City	Affected buildings, %
Norilsk	10%
Tiksi	22%
Dudinka	55%
Dikson	35%
Pevek, Amderma	50%
Chita	60%
Vorkuta	80%

Anisimov and Lavrov, 2004



Photo by V. Romanovsky



DT_s	Decrease of bearing capacity of foundations, % from normal	
	Under buildings	Under pipelines
0.5	5	10
1.0	15	20
1.5	20	30
2.0	50	60

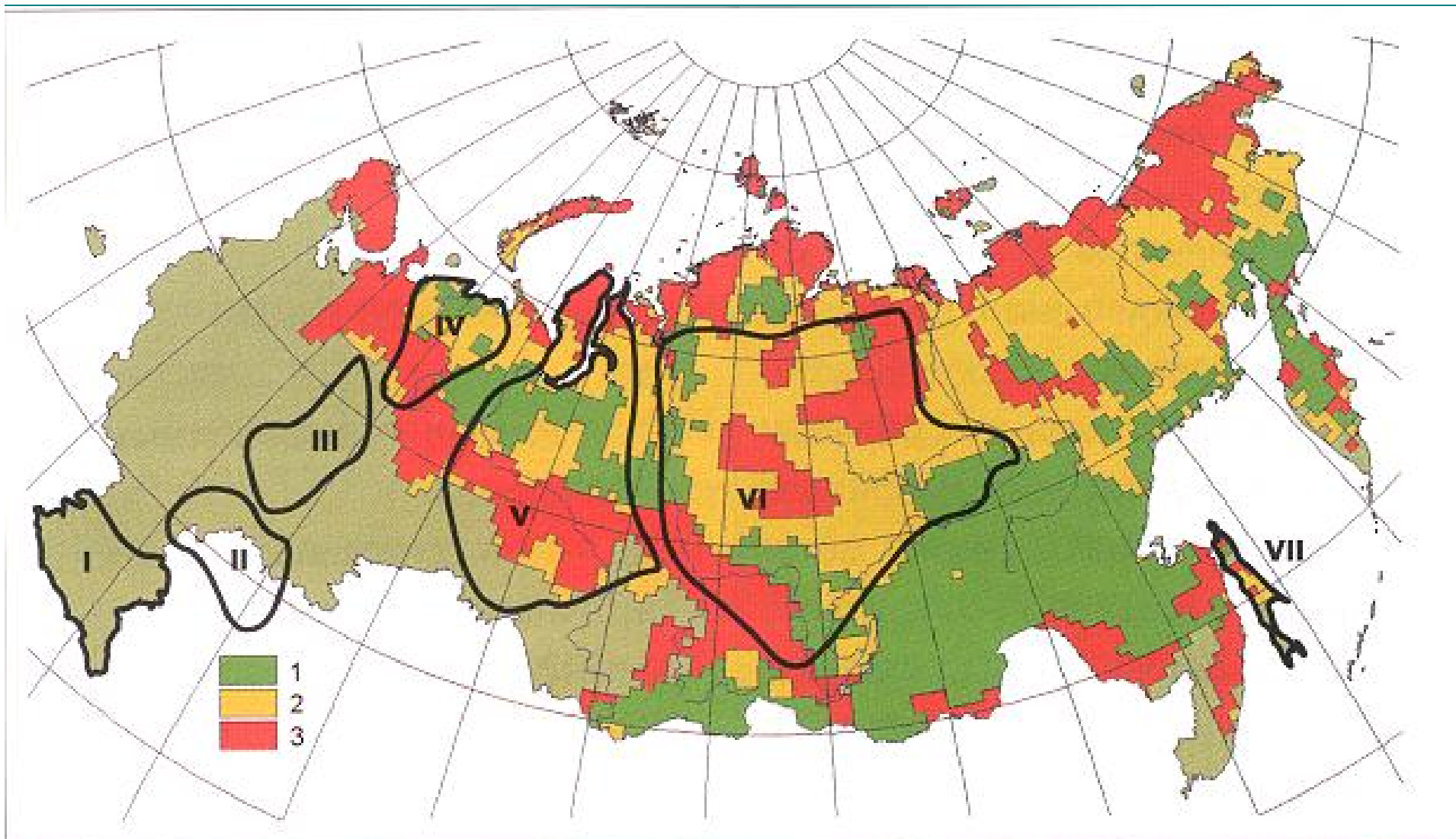
Khrustalev, 2001

Permafrost hazard index, I_g

(Anisimov and Lavrov, 2004)

$$I_g = k \cdot (1 + S) \cdot (z_2 - z_1) \cdot C_w / z_1$$

- z_i - maximum summer thaw depth under the modern (index 1) and projected for the future (index 2) climate,
- C_w - volumetric ground ice content,
- S - salinity of soil,
- K - constant scaling factor.



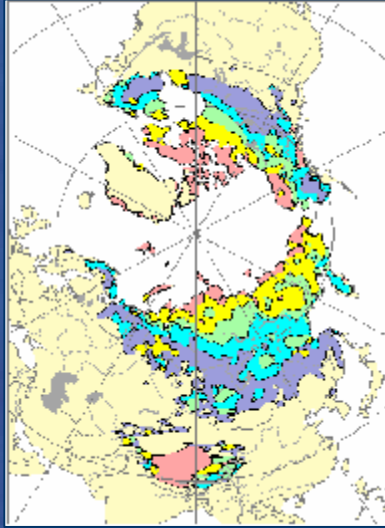
Regions with low (1), moderate (2) and high (3) susceptibility to permafrost hazard, GFDL climatic projection for mid-21st century.

Numbered regions on the map designate provinces of oil and gas exploration.

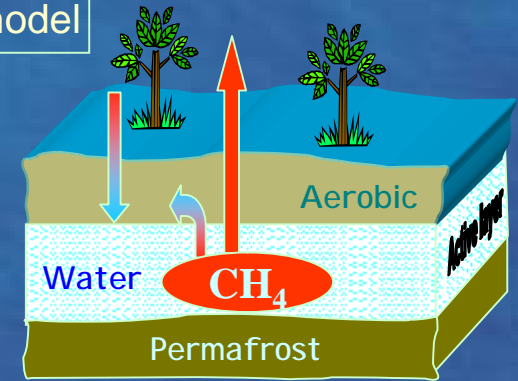
(Anisimov and Lavrov, 2004)

Feedback to the global climate system through methane emission from Russian Arctic wetlands.

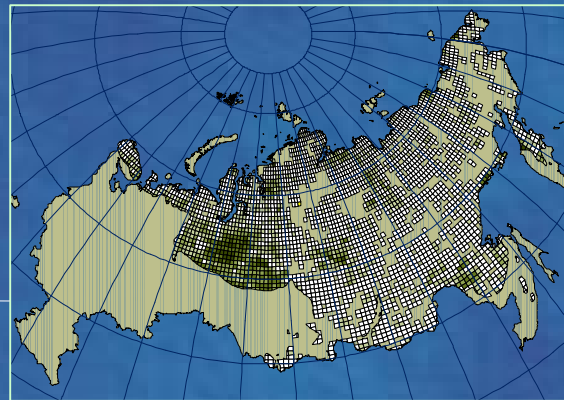
Permafrost model



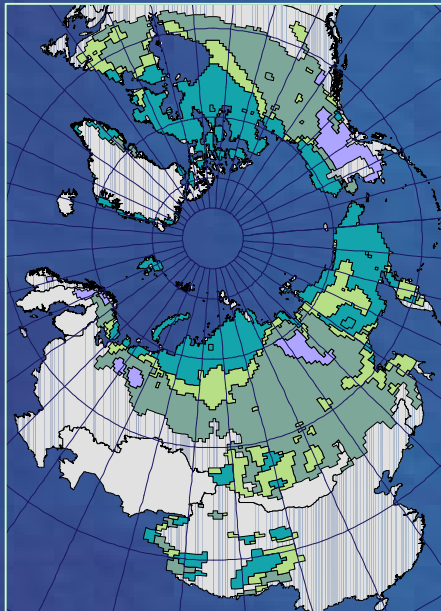
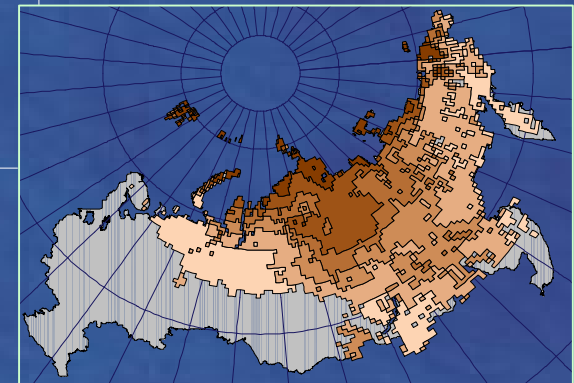
Methane model



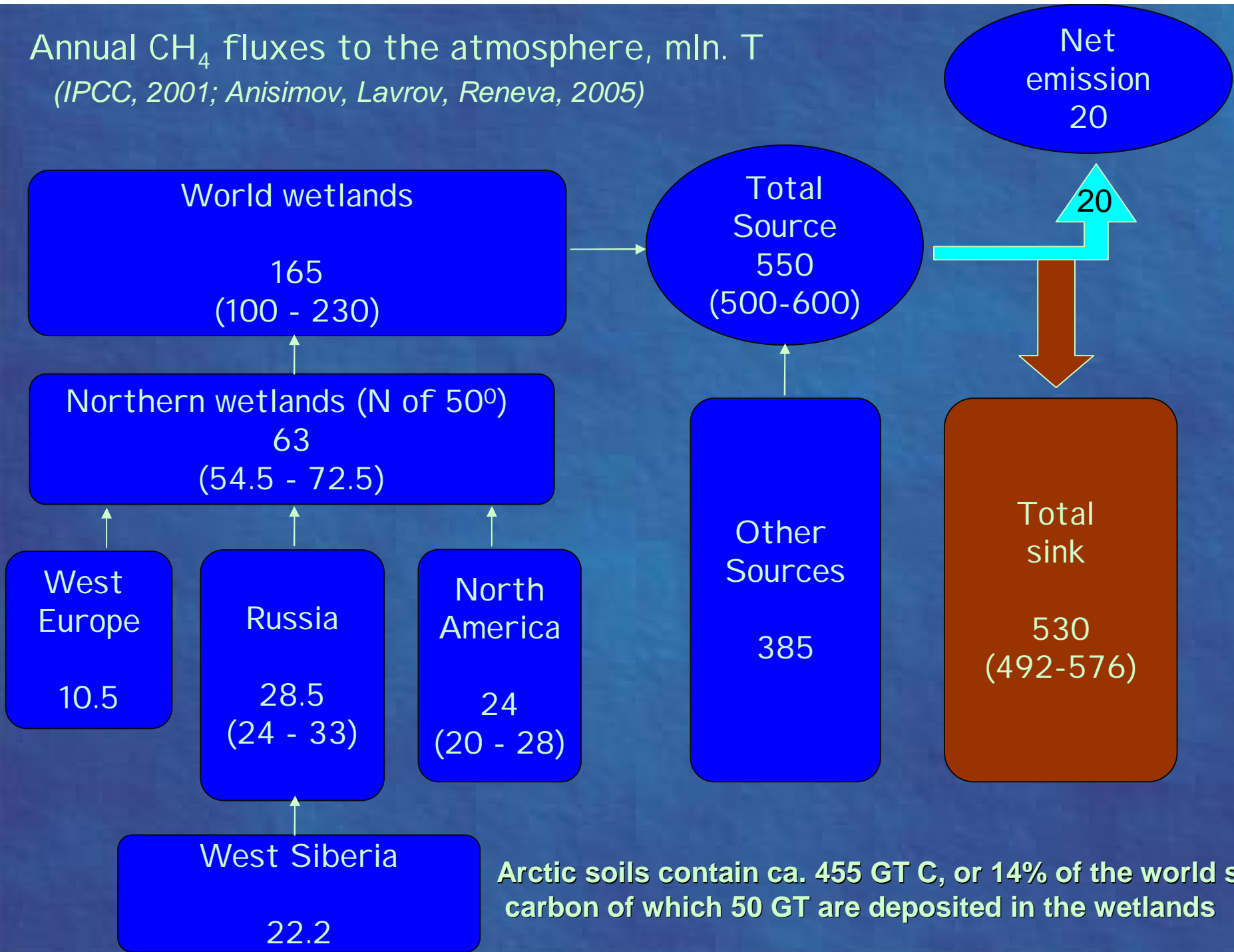
Area of wetlands



Methane emission

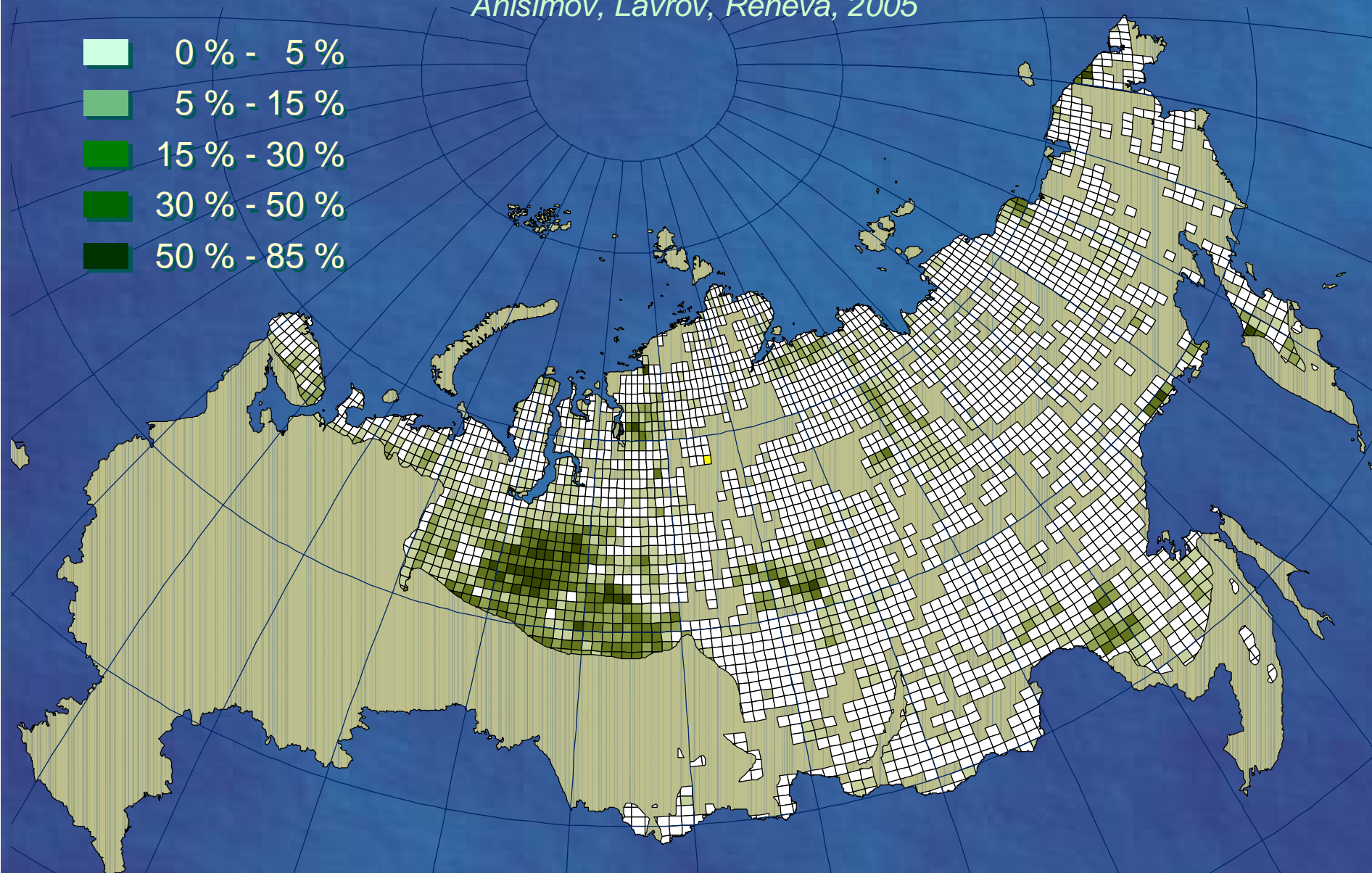
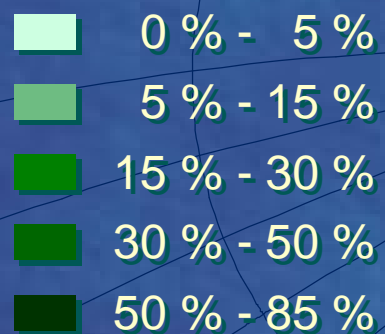


Annual CH₄ fluxes to the atmosphere, mln. T
(IPCC, 2001; Anisimov, Lavrov, Reneva, 2005)



Fraction of land occupied by wetlands in Russia

Anisimov, Lavrov, Reneva, 2005



area of bogs – 0,76 mln. km², permafrost bogs – 0,35 mln. km²

Projected changes of the volume of seasonally thawing peat (km³). Water content 600 мм/м / 800 мм/м.

(Anisimov, Lavrov, Reneva, 2005)

years models	2025	2050	2080
NCAR	-	80,61 / 88,59	-
GFDL	80,88 / 90,87	137,82 / 150,40	153,94 / 173,02
ECHAM – 4	-	228,92 / 249,04	-

Spasskaya Pad research station near Yakutsk





Parameterization of methane fluxes:

Flux / thaw depth ratio:

$$J_1 / J_0 = \sqrt{H_{d1} / H_{d0}}$$

Flux / ground temperature ratio:

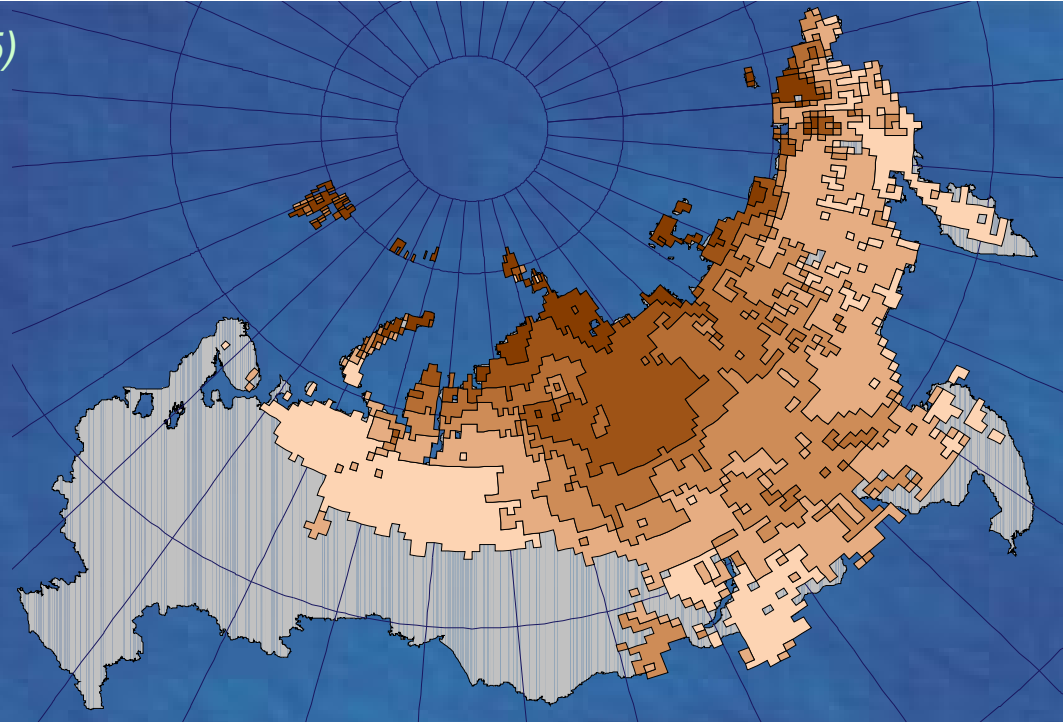
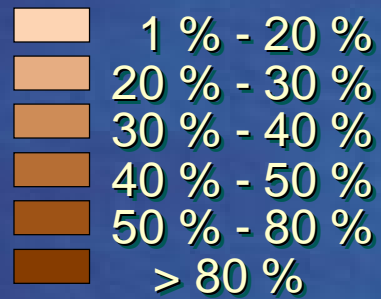
$$J_1 / J_0 = \exp 0.1(T_1 - T_0)$$

Combined equation:

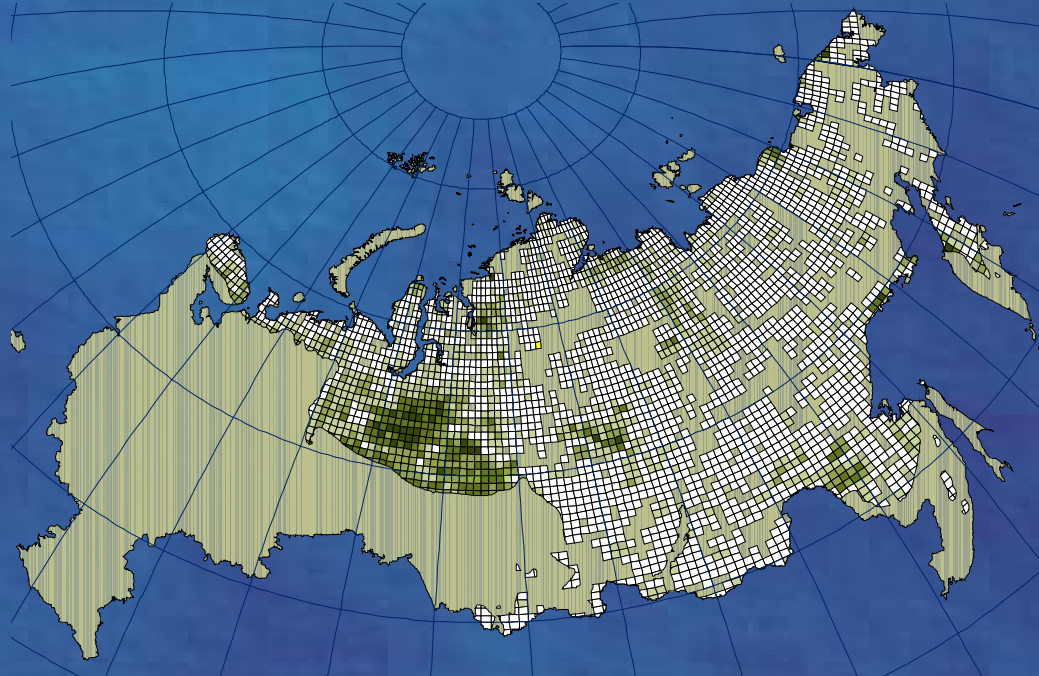
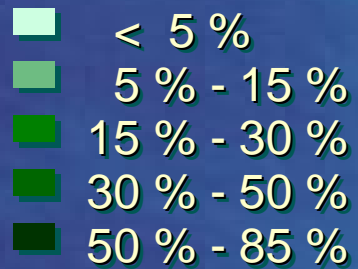
$$J_1 / J_0 = \exp 0.1(T_1 - T_0) \sqrt{H_{d1} / H_{d0}}$$

(Anisimov, Lavrov, Reneva, 2005)

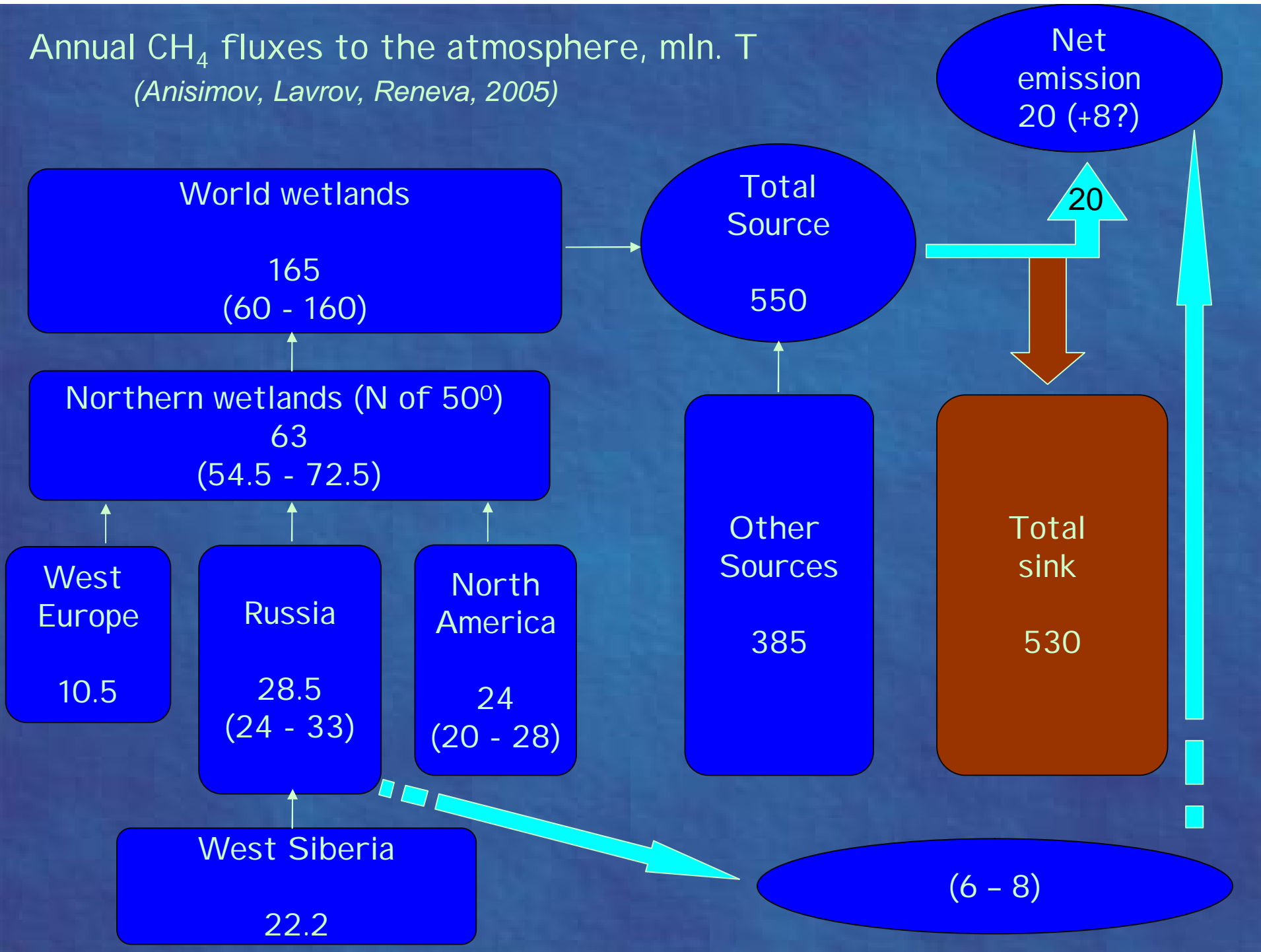
Predicted changes of methane flux, GFDL scenario for 250



Fraction of permafrost occupied by frozen wetlands

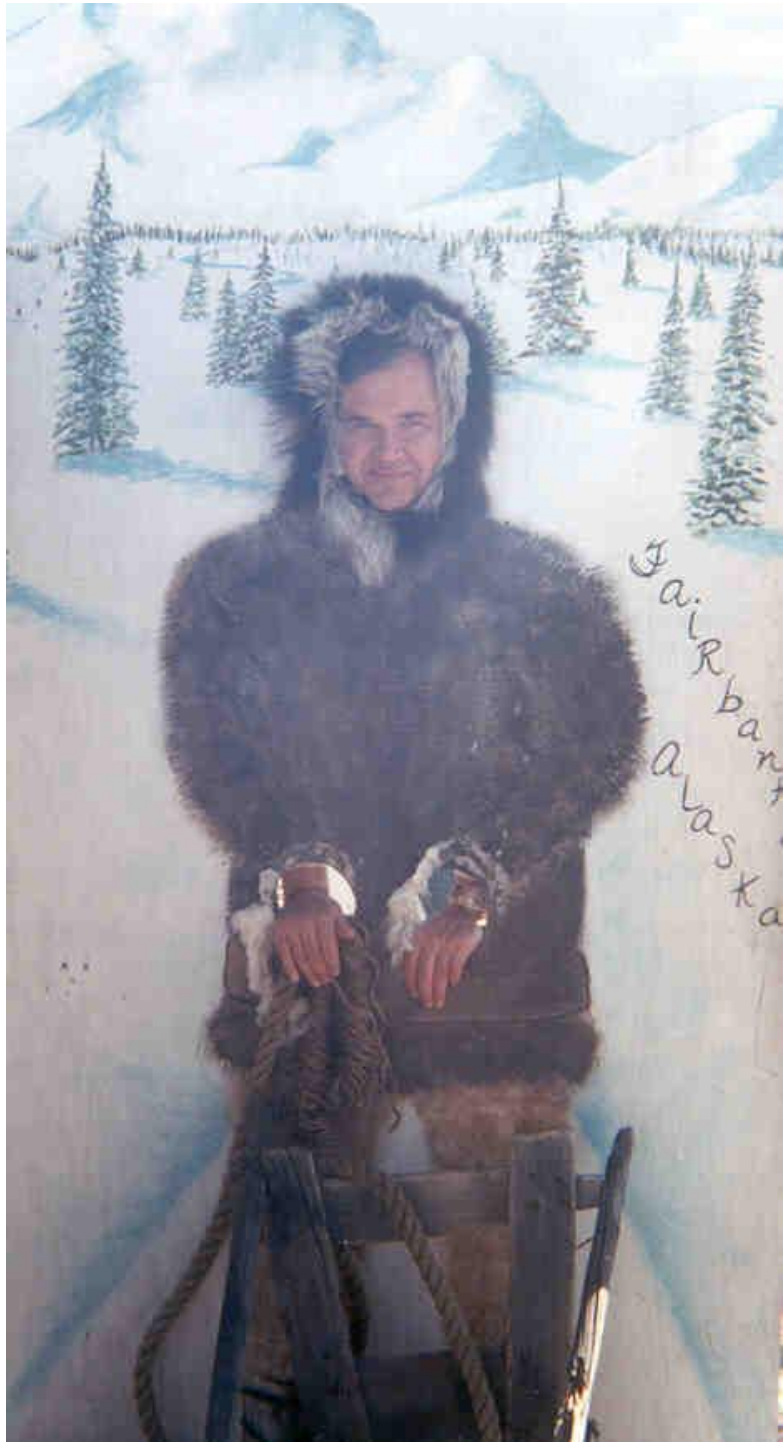


Annual CH₄ fluxes to the atmosphere, mln. T
(Anisimov, Lavrov, Reneva, 2005)



Projected changes of CH₄ emissions from wetlands

CH ₄ changes	Climate change	Source
+25%	+1.5 °C	Anisimov et al., 2005
156-277mln.t./yr (+78%)	CO ₂ doubling in GISS GCM (+3.4 °C)	Schindell et al., 2004
±20% change	Uniform ±1 °C temperature change	Walter et al., 2001
±8% change	±20% precipitation change	Walter et al., 2001
26% increase	global warming of 2.5 °C	Liu, 1996; Prinn et al, 1999
+19% increase	Uniform +2 °C temperature change	Cao et al., 1998
+21% increase, N60+	+2 °C temperature and +10% precipitation change	Cao et al., 1998
Decreased emission	> 4 °C temperature change	Cao et al., 1998
+38%, N60+	+10 cm thaw depth	Zhuang et al., 2004
0 to +25%, N60+	+10 cm water table rise	Walter and Heimann, 2000



Thank you !